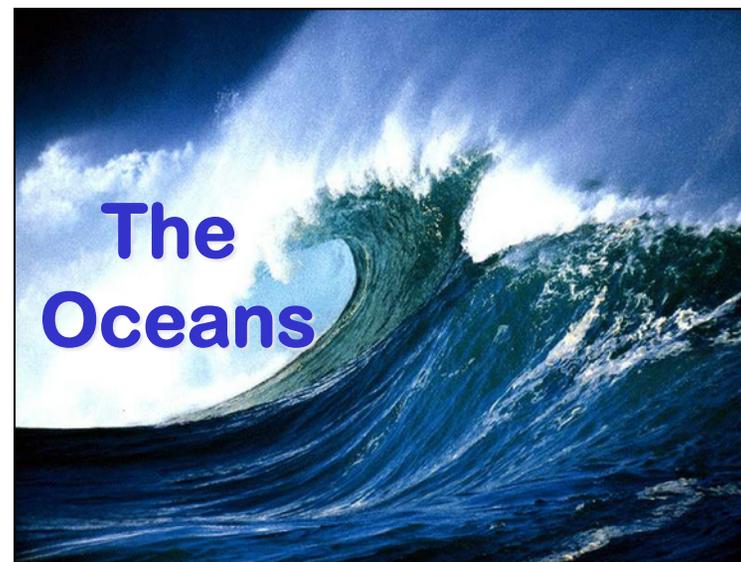
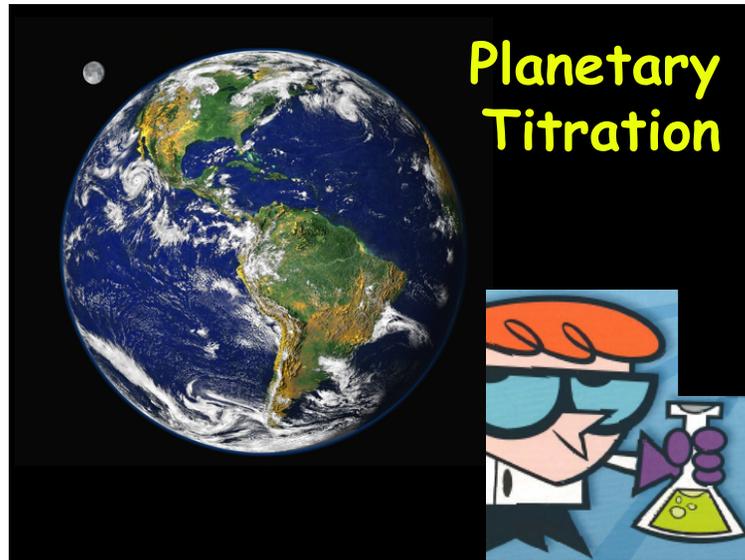


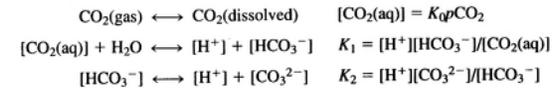
- ### Where Has All the Carbon Gone?
- Into the **oceans**
 - **Solubility pump** (CO₂ very soluble in cold water, but rates are limited by slow physical mixing)
 - **Biological pump** (slow "rain" of organic debris)
 - Into the **land**
 - **CO₂ Fertilization** (plants eat CO₂ ... is more better?)
 - **Nutrient fertilization** (N-deposition and fertilizers)
 - **Land-use change** (forest regrowth, fire suppression, woody encroachment ... but what about Wal-Marts?)
 - Response to **changing climate** (e.g., Boreal warming)



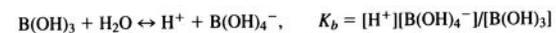
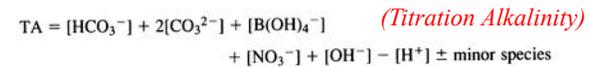


Carbonate Equilibria in Solution

Three equations (equilibria) in five unknowns



Add two more constraints

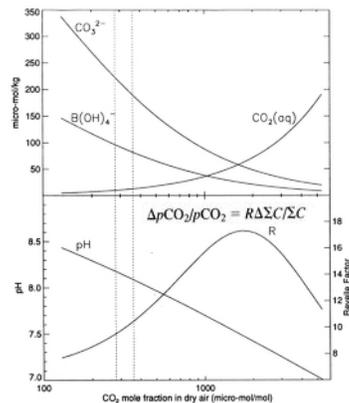


(Boric acid dissociation)

$$\Sigma B = 1.179 \times 10^{-5} S \text{ mol/kg}$$

(Salinity)

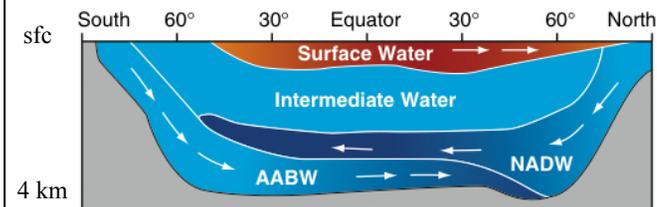
The "Revelle Factor"



- As CO₂ is added to the atmosphere, CO₃²⁻ is consumed to form additional HCO₃⁻
- As CO₃²⁻ becomes more scarce, Revelle factor increases, severely limiting capacity of ocean to absorb more CO₂
- At very high pCO₂, pH decreases enough that sedimentary CaCO₃ dissolves, providing CO₃²⁻ and decreasing R again

Figure 12.2. The response of surface ocean waters at 16°C to increasing atmospheric concentrations of CO₂, which is plotted on a logarithmic scale. The dissolved CO₂(aq) increase is directly proportional to pCO₂ (Henry's Law). The decreases of CO₃²⁻ and B(OH)₄⁻ have been calculated with the constants of Table 12.1. In the lower panel, pH is referred to the left axis, and the Revelle factor, R, (see main text) is referred to the right axis. The vertical dashed lines indicate the preindustrial (280 ppm) and the 1995 level of CO₂ (360 ppm).

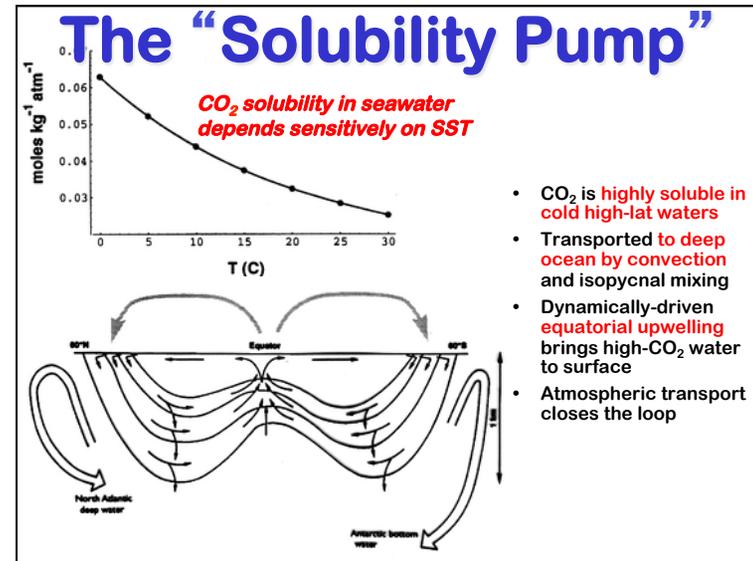
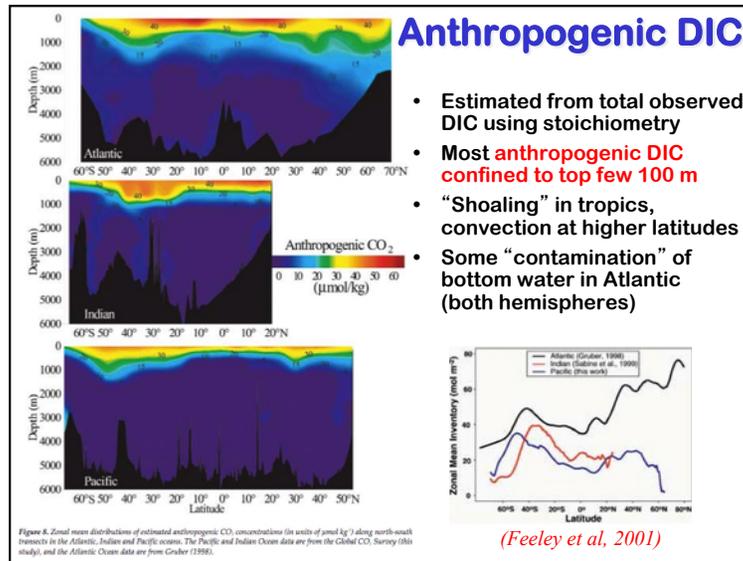
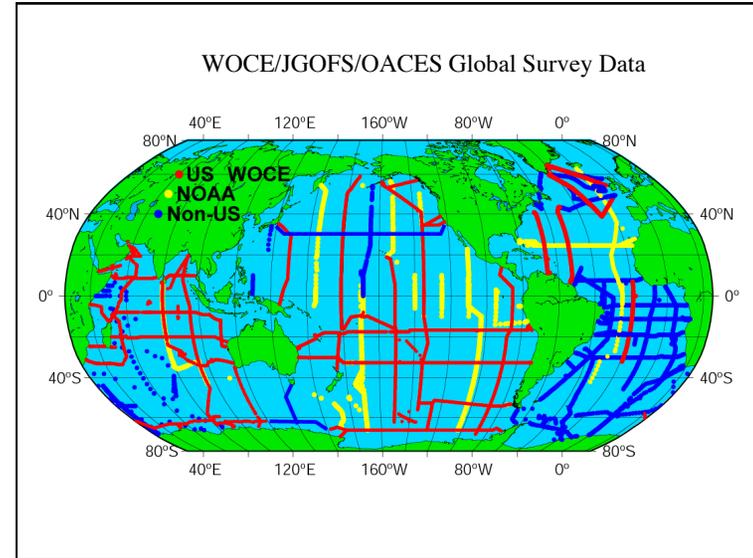
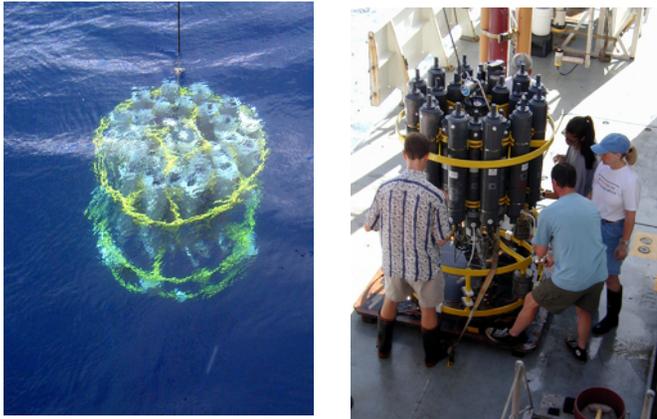
Vertical Structure of the Oceans

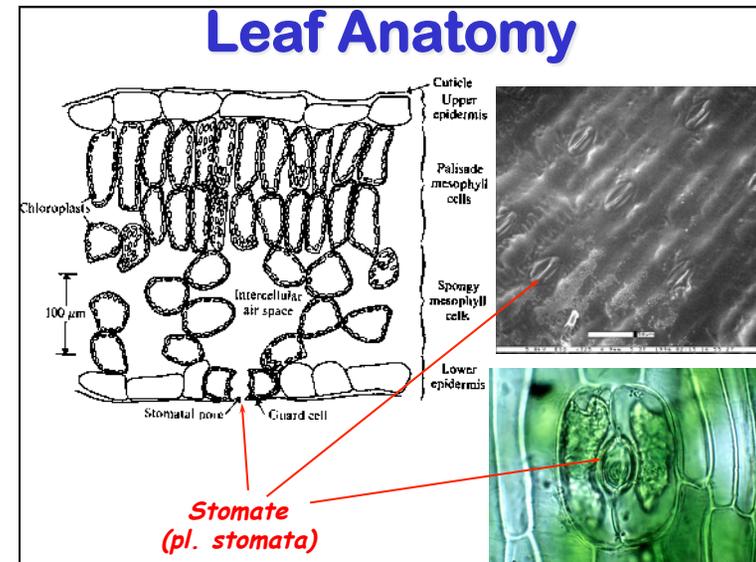
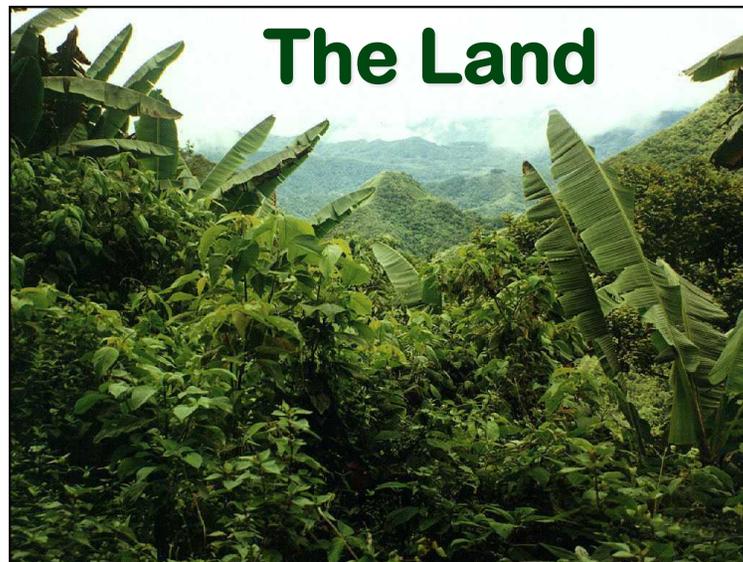
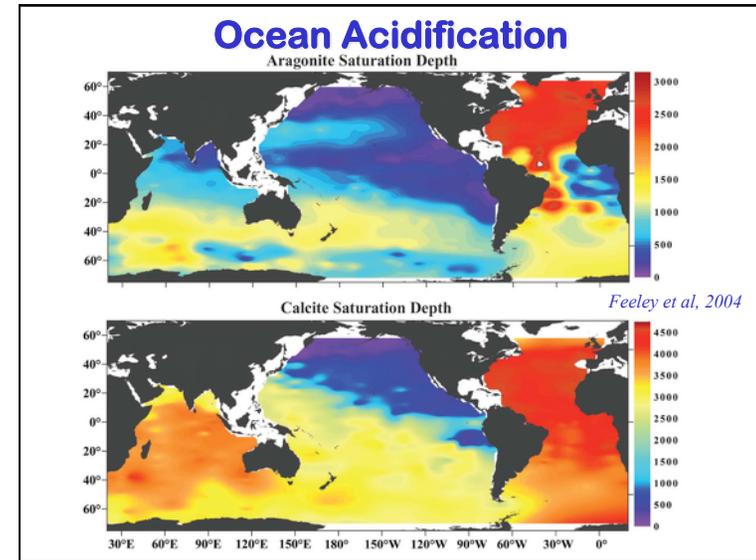
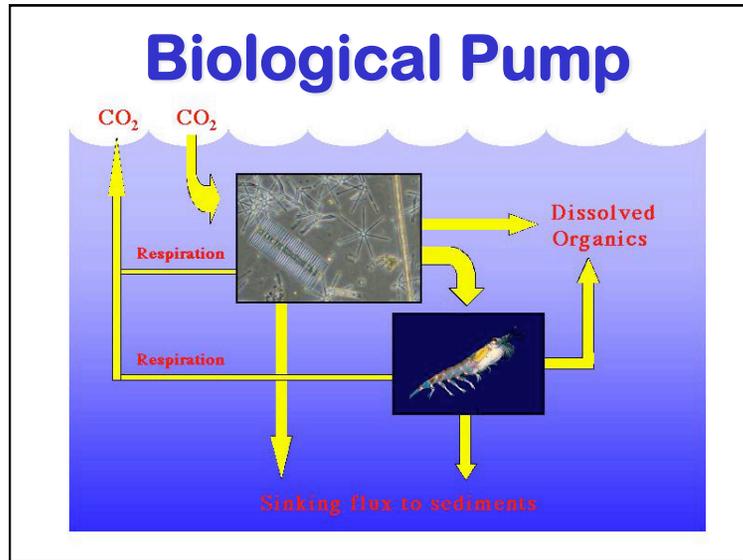


- Increased nutrients & dissolved CO₂
- Warm, low nutrients, & oxygenated

- Warm buoyant "raft" floats at surface
- Cold deep water is only "formed" at high latitudes
- Very stable, hard to mix, takes ~ 1000 years!
- Icy cold, inky black, most of the ocean doesn't know we're here yet!

Observing the Deep Ocean





Carbon and Water

- Plants eat CO₂ for a living
- They open their stomata to let CO₂ in
- Water gets out as an (unfortunate?) consequence
- For every CO₂ molecule fixed about 400 H₂O molecules are lost

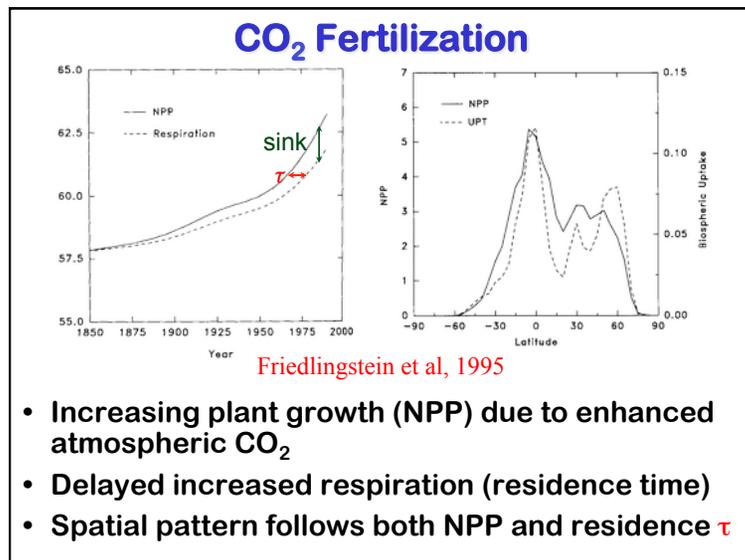
Canopy Carbon Balance

Carbon Storage & Turnover in Land Ecosystems

- Photosynthesis converts inorganic gas into living plants
- Plants die and become "litter"
- Microbes eat litter and poop soil carbon
- Soil carbon is eventually also eaten by microbes, but some lasts for many centuries

Net Primary Production and mean residence times of carbon determines source/sink potential of ecosystems

Where r is the fractional perturbation to NPP or MRT with units (1/time)



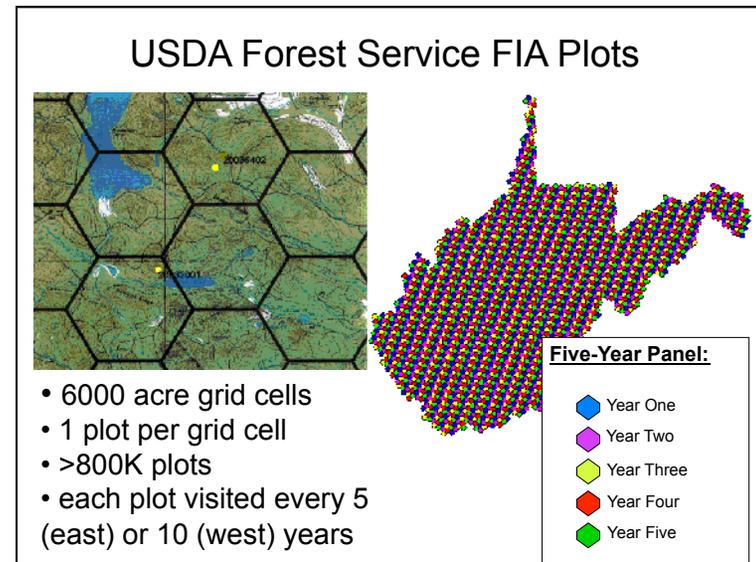
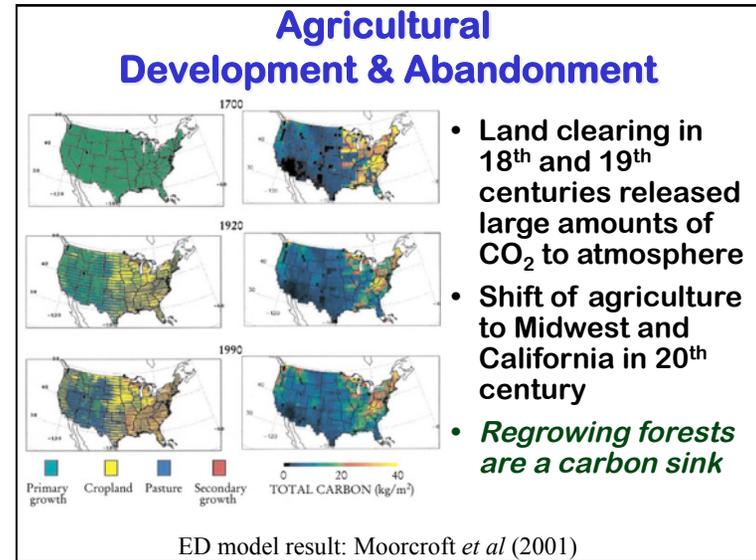
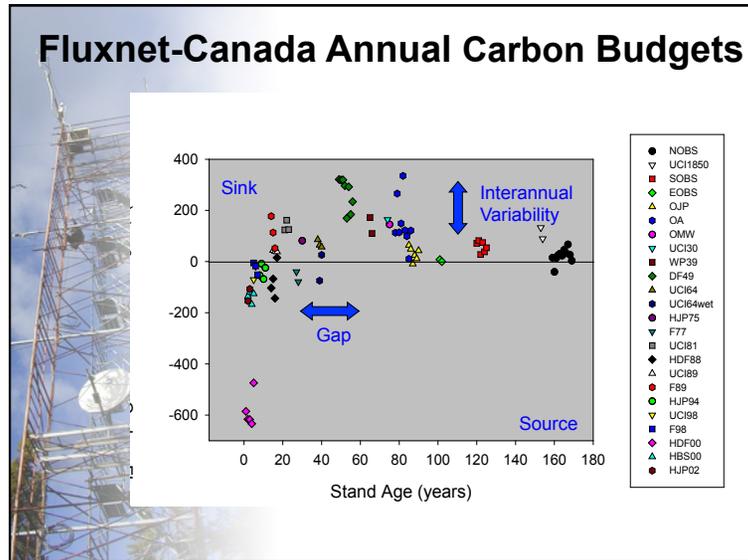
Free Air Carbon Enrichment (FACE)

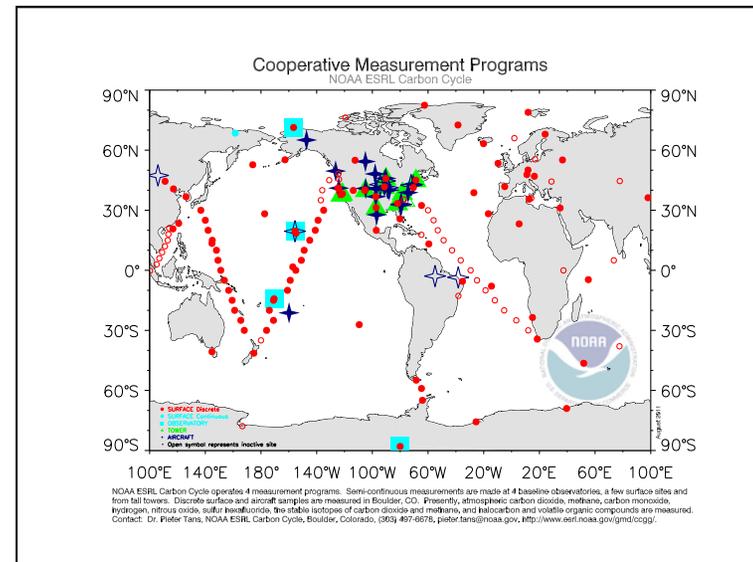
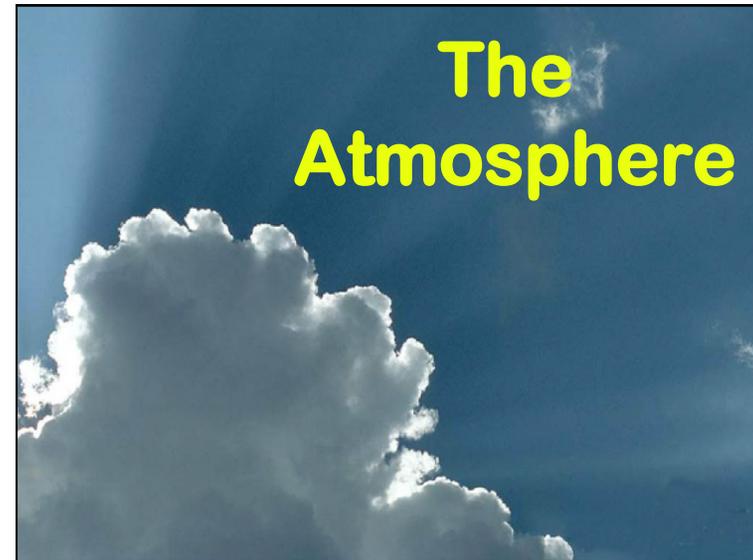
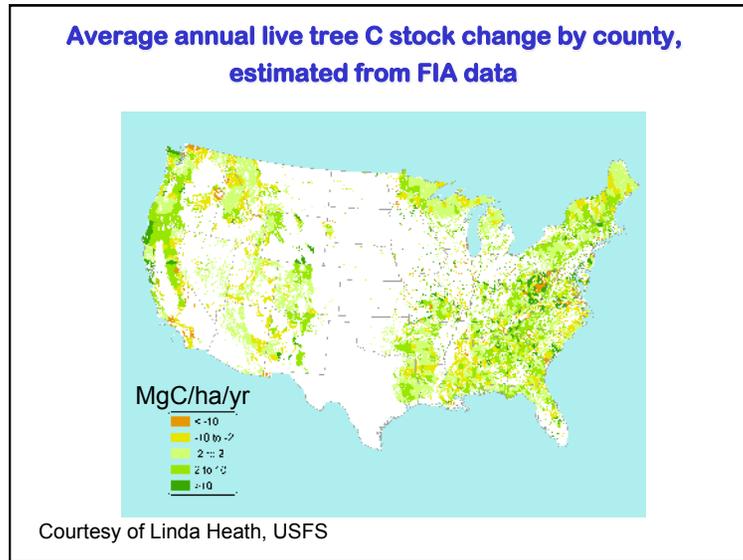
- Fumigation rings maintain steady levels of **elevated CO₂** in canopies under changing weather conditions
- Control and replicated treatments test effects of CO₂, water, N, etc

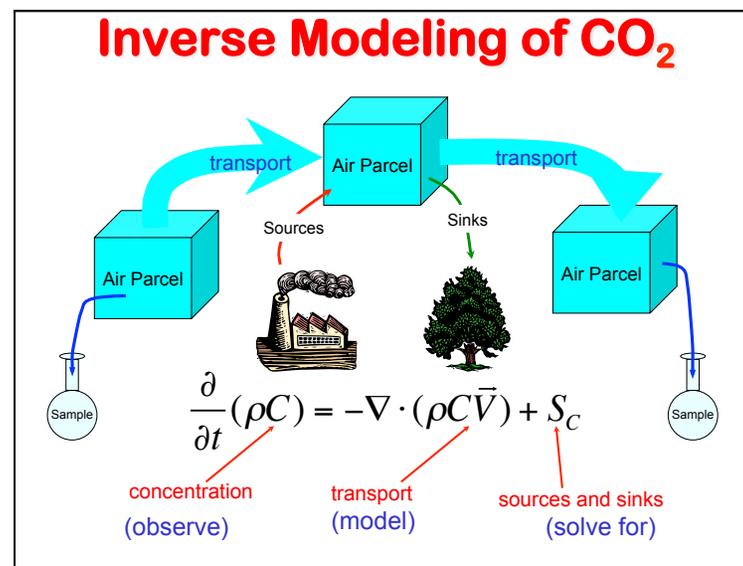
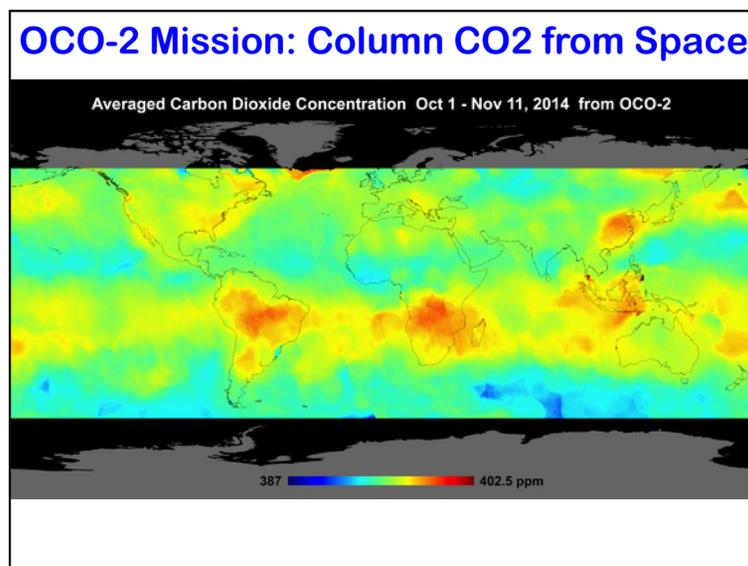
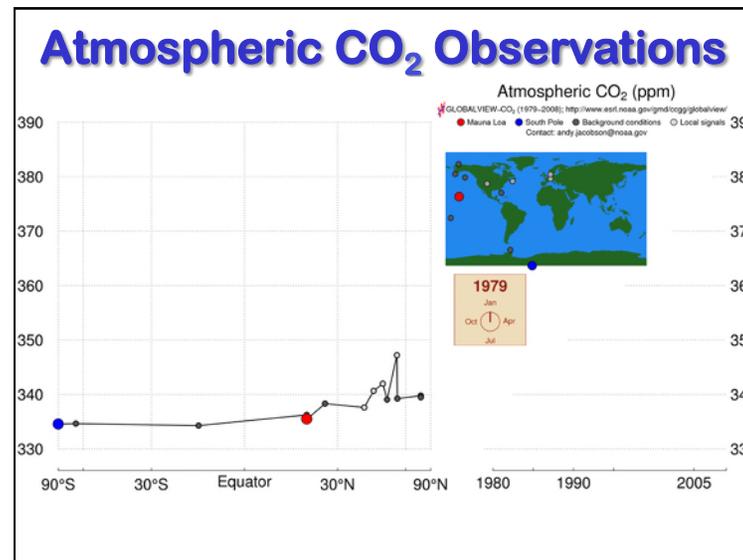
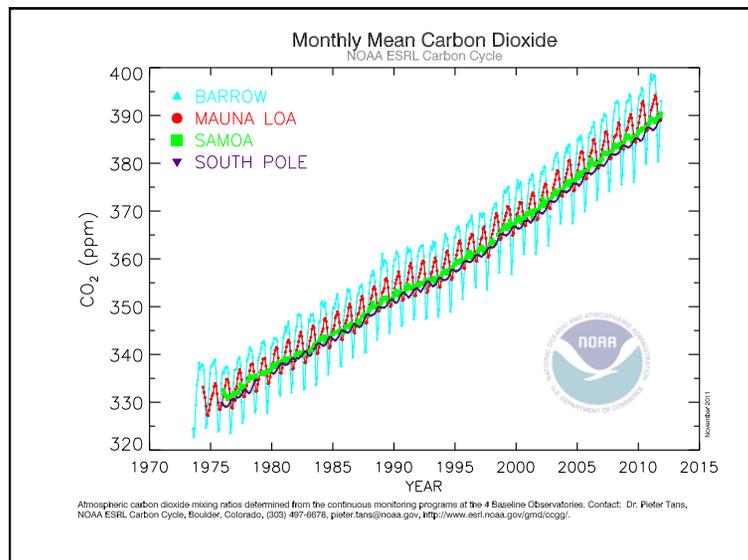


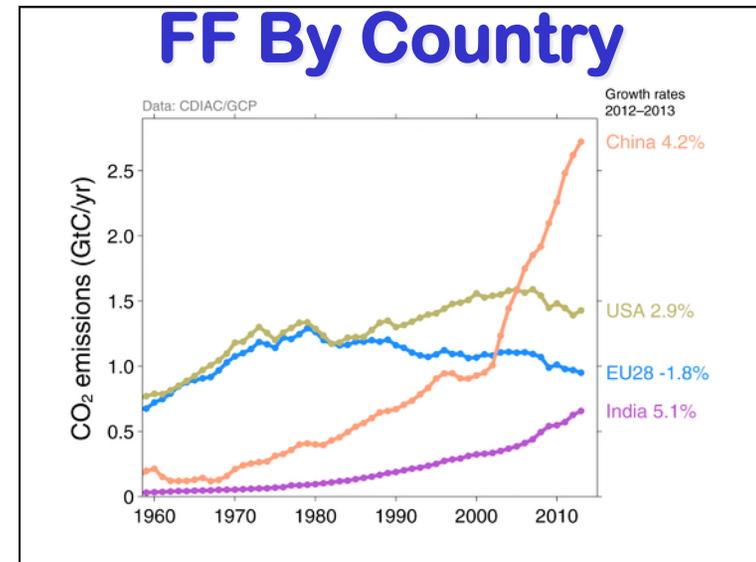
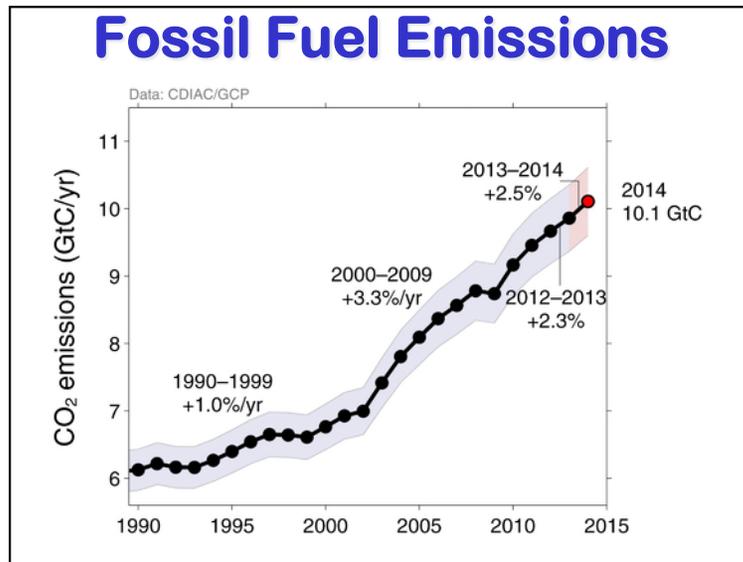
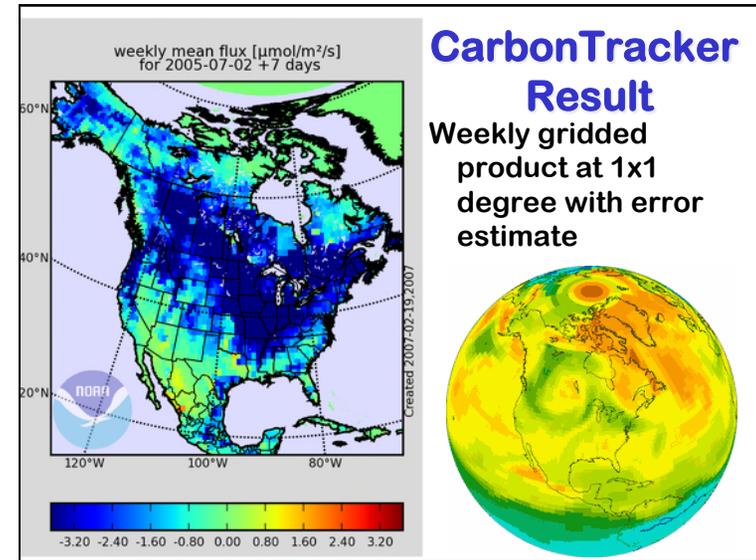
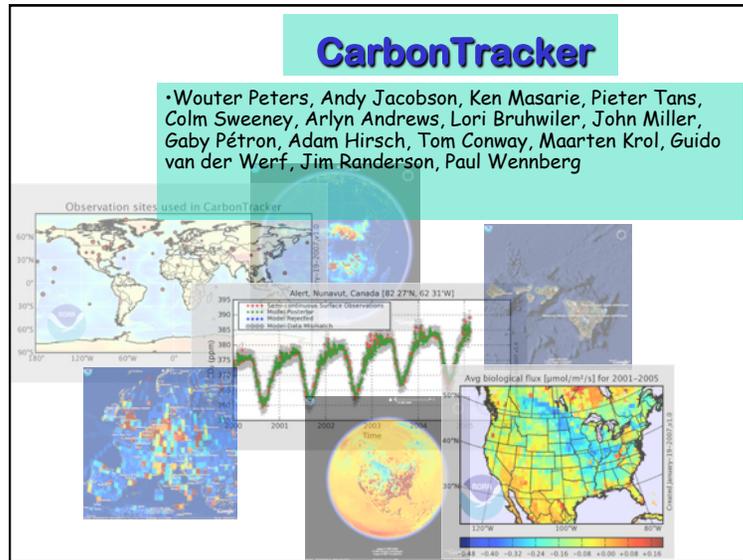
Disturbance and Recovery

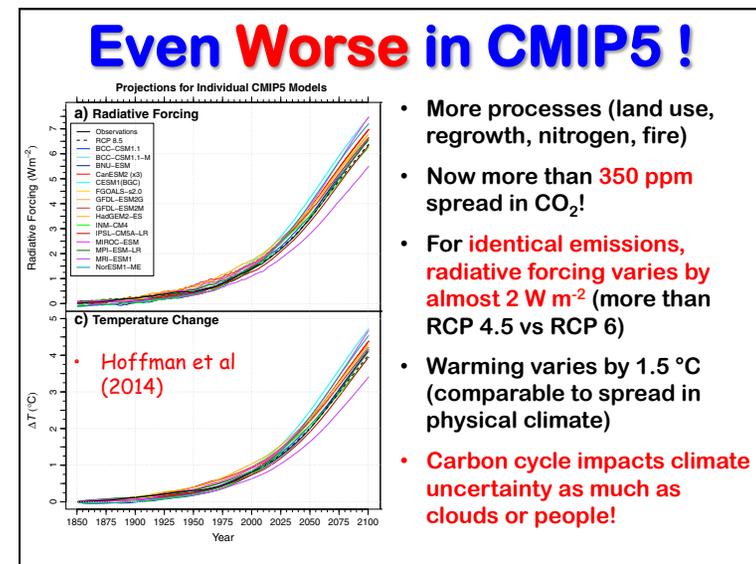
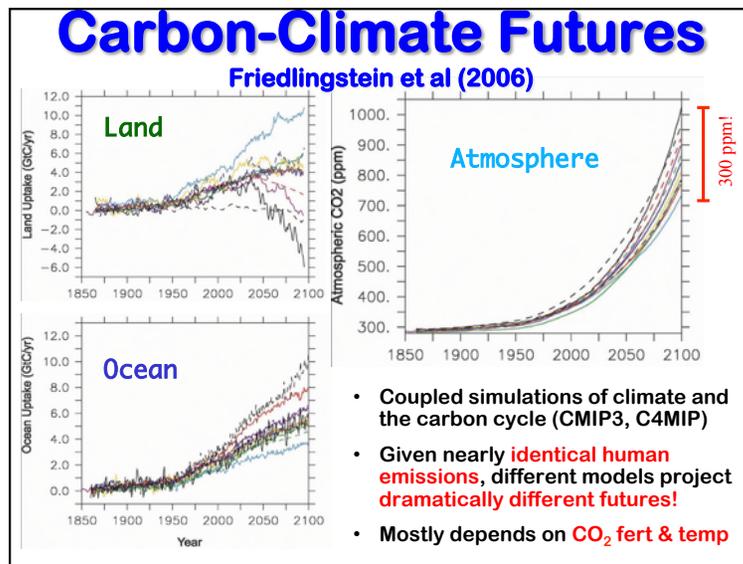
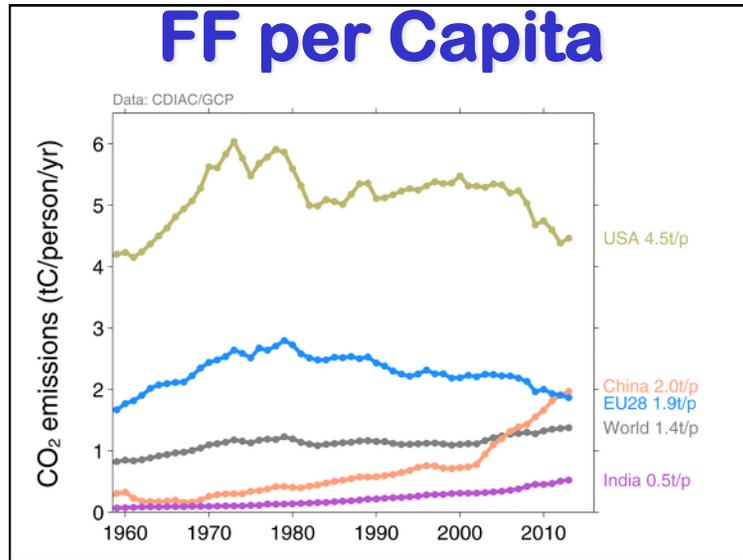
Planted	2000	1988	1949
Height (m)	0.8	3-8	30-35





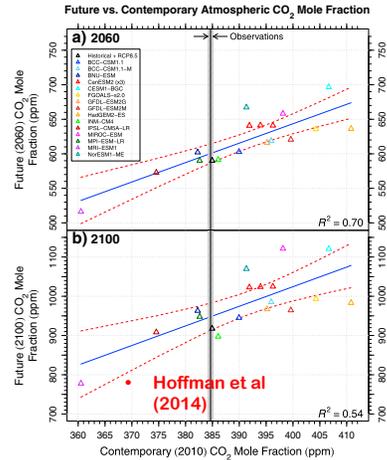




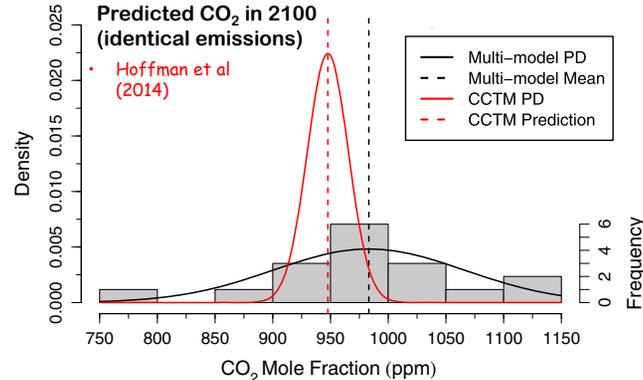


Past as Prelude

- Models that underpredict contemporary CO₂ also predict low CO₂ in the future, and vice versa
- Evaluation of past carbon cycle simulations constrain future feedback



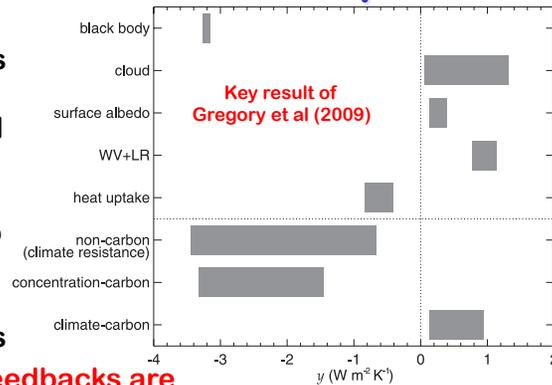
Carbon Constraint



- Fivefold reduction in model spread in 2100
- No mechanism ... simple scalar multiplication of sinks

Climate Feedback Comparison

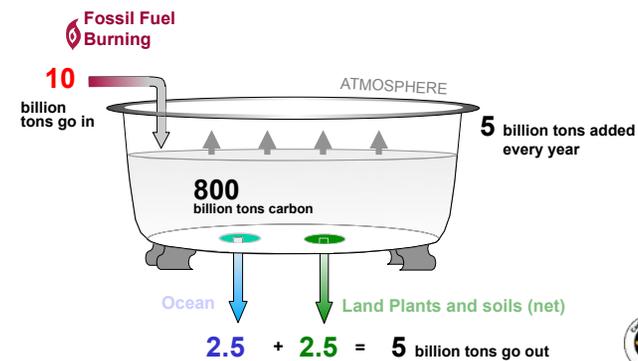
- Carbon feedbacks are compared “apples-and-apples” to other climate feedbacks



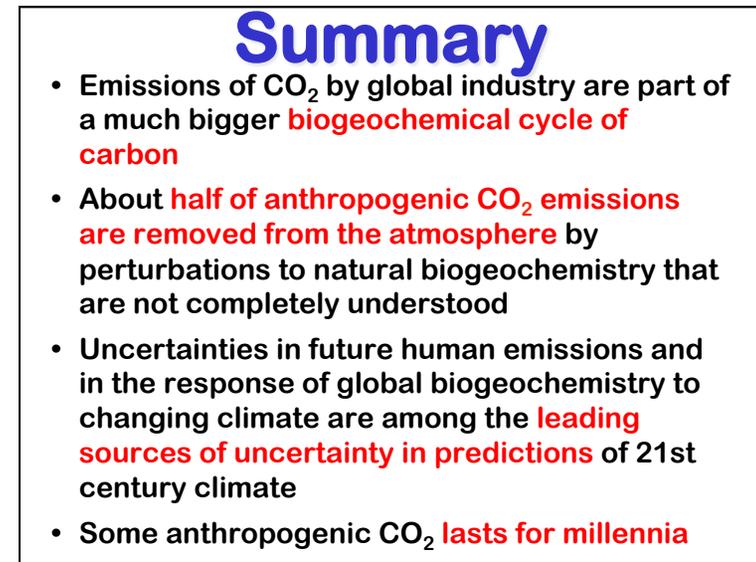
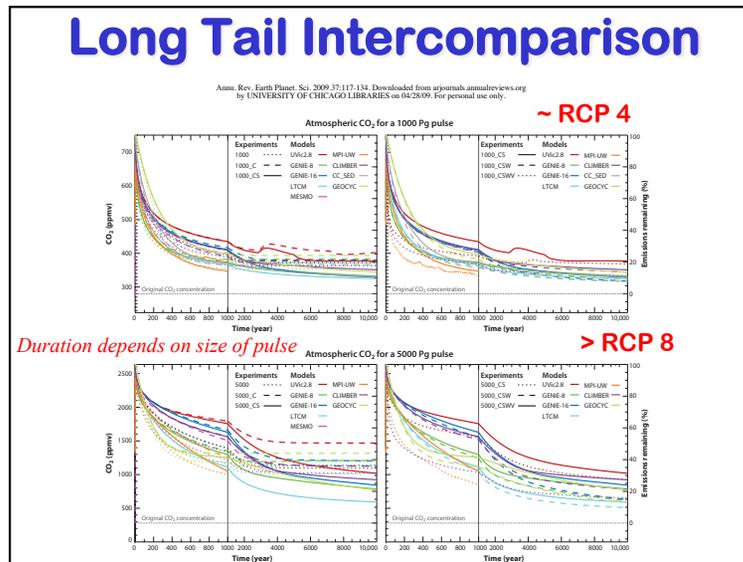
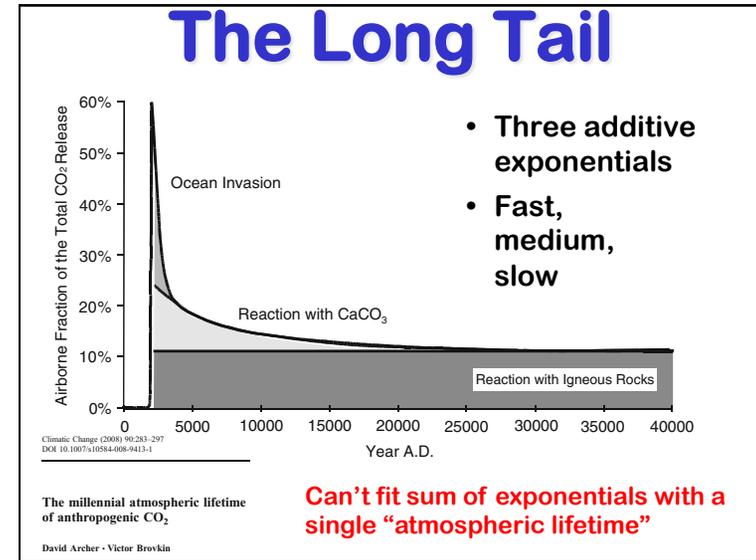
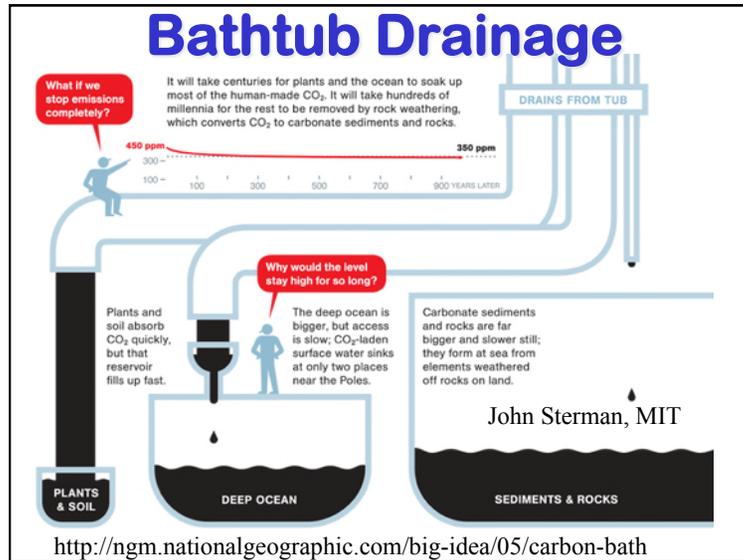
- Carbon feedbacks are strongest and most uncertain of all feedback processes in climate system

FIG. 2. (top) A comparison of components of climate feedback and ocean heat uptake, and (bottom) a comparison of the combined oceanic response (the climate resistance; the sum of the terms in the top part) with the carbon cycle feedbacks for forcing resulting from CO₂ emissions. The blackbody, cloud, surface albedo, and WV + LR (water vapor and longwave) terms are from Soden and Held (2006). The heat uptake term is the ocean heat uptake efficiency as from Gregory and Foster (2009) and the climate resistance is that λ evaluated from CMIP3 AOGCMs, with its uncertainty amplified by the concentration-carbon feedback as described in the text. The carbon cycle contributions are as calculated in this paper from the CMIP3 models of Friedling et al. (2006). We show climate feedback parameters as $-\lambda$, so that positive terms lead to increased climate warming for a positive forcing; the blackbody response is shown as a negative climate feedback $-\lambda_{bb}$, and the ocean heat uptake efficiency and climate resistance are likewise negative terms. The bars indicate approximate 95%–95% confidence intervals (mean $\pm 1.65 \times SD$).

CO₂ “Budget” of the World



Climate forcing comes from the water, not from the faucet!



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