

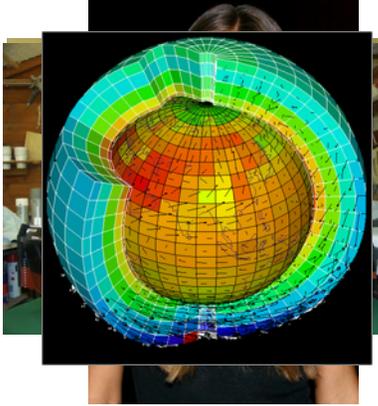
The Layer Model of the Greenhouse Effect

ATS 150
Lecture 4

Please read Chapter 3
in Archer Textbook

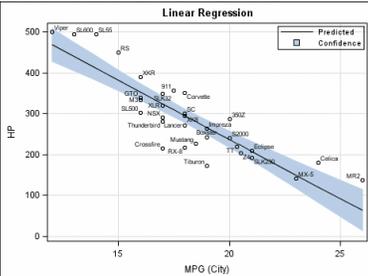
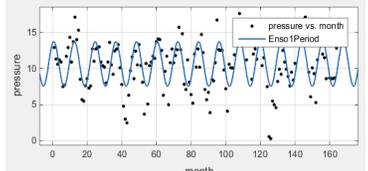
What is a Model?

- Models represent reality in a simplified or idealized way
- Used to understand or predict
- Doesn't have to be realistic to be useful

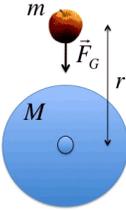


Empirical Models

- Generalized mathematical formulation with adjustable coefficients
- Combinations of
 - polynomials
 - exponential growth & decay
 - Periodic sines and cosines
- Coefficients fit to data (e.g., least squares)
- Interpolation or extrapolation

Deterministic Models



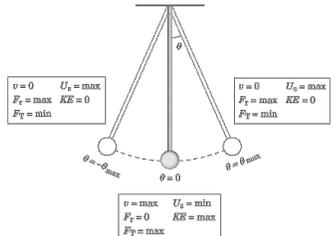
$$F = ma$$

$$\frac{GMm}{r^2} = ma$$

Then, cancelling m on both sides:

$$a = \frac{GM}{r^2} = g$$

- Formulated as "cause and effect"
- Common in physics and chemistry
- Usually take the form of differential equations
- Initial & boundary-value problems
- May still have adjustable coefficients



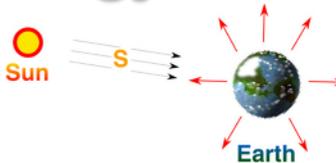
Energy Balance



Energy In = Energy Out

- Let the rate of energy flow from the Sun to the Earth be called F_{in}
- Let the rate of energy flow from the Earth to outer space be called F_{out}

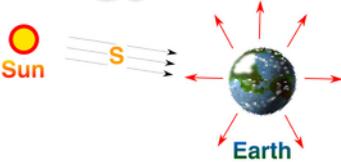
Energy Balance



$F_{in} = F_{out}$

- Assume both Sun and Earth are blackbodies ($\epsilon=1$), so $F = \sigma T^4$
- F_{in} = absorbed sunlight x daylight area
- $F_{out} = \sigma T_{earth}^4$ x total area

Energy Balance

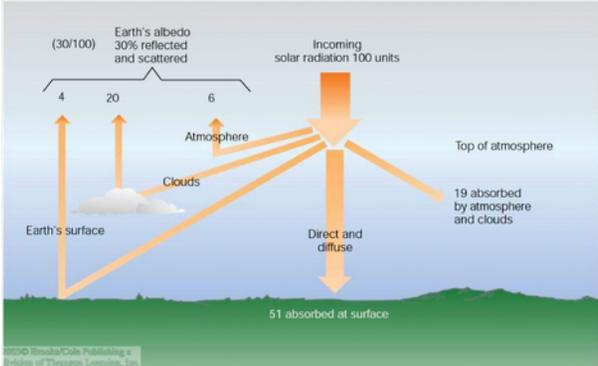


$F_{in} = F_{out}$

F_{in} = Solar brightness x (1 - albedo) x (area of Earth's shadow)

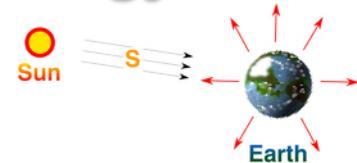
$F_{out} = \sigma (T_{earth})^4$ x (area of Earth's surface)

Solar Radiation



- 30% reflected by clouds, air, dust, and surface
- 19% absorbed by the atmosphere (mostly clouds)
- 51% absorbed at the surface

Energy Balance

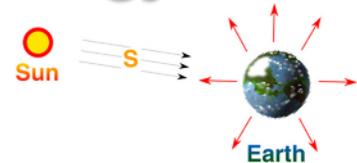


$F_{in} = F_{out}$

$F_{in} = \text{Solar brightness} \times (1 - \alpha) \times \pi r^2$

$F_{out} = \sigma (T_e)^4 \times 4\pi r^2$

Energy Balance



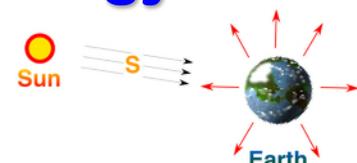
$F_{in} = F_{out}$

$$S(1 - \alpha)\pi r^2 = 4\pi r^2 \sigma T^4$$

$$S(1 - \alpha) = 4\sigma T^4$$

$$\frac{S(1 - \alpha)}{4\sigma} = T^4$$

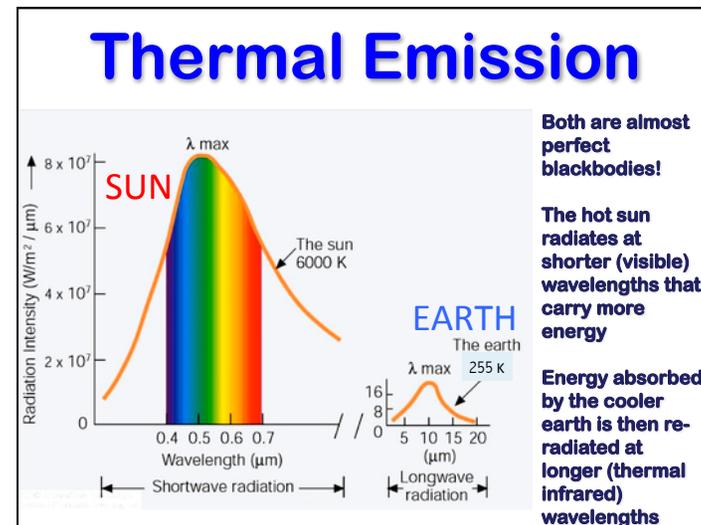
Energy Balance



$F_{in} = F_{out}$

$$T = \left(\frac{S(1 - \alpha)}{4\sigma} \right)^{1/4} = \left(\frac{(1360 \text{ Wm}^{-2})(1 - 0.3)}{4(5.67 \times 10^{-8} \text{ Wm}^{-2} \text{ K}^{-4})} \right)^{1/4} = 255 \text{ K}$$

But the observed surface temperature is about 288 K



Works for Other Planets Too

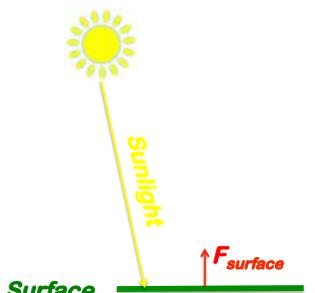


$F_{in} = F_{out}$

$$\frac{S(1-\alpha)}{4\sigma} = T^4$$

Experiment! <http://tinyurl.com/planetary-balance>

Bare Rock Model

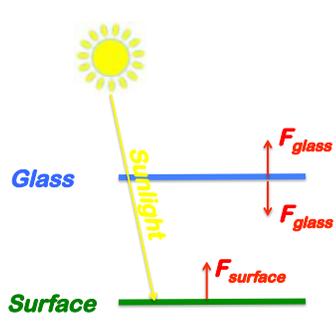


$F_{in} = F_{out}$

$$\frac{S(1-\alpha)}{4} = \sigma T^4$$

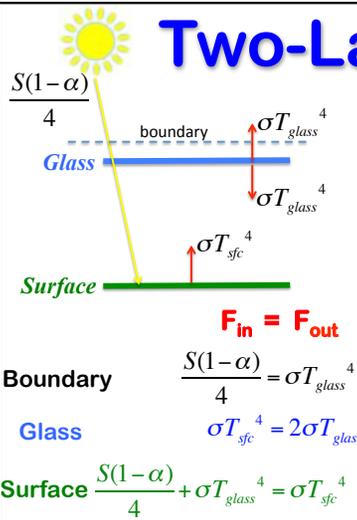
- If sunlight were the only source of energy, the surface of the Earth would be **way too cold for life!**
- 255 °K = -18 °C = -1 °F

Two-Layer Model



- Now imagine a layer of glass above the solid rock
- Glass is **transparent to sunlight**
- Glass is **opaque to "Earthlight"** (thermal IR)

Two-Layer Model



- Sun shines right through transparent glass layer
- Glass layer emits both upward and downward
- Glass layer "feels" both space and surface
- Surface "feels" both Sun and glass

$F_{in} = F_{out}$

Boundary $\frac{S(1-\alpha)}{4} = \sigma T_{glass}^4$

Glass $\sigma T_{sfc}^4 = 2\sigma T_{glass}^4$

Surface $\frac{S(1-\alpha)}{4} + \sigma T_{glass}^4 = \sigma T_{sfc}^4$

Seen from Above

Both models:

- Absorb the **same** sunshine
- Reflect the **same** amount of sunlight back out
- Emit the **same** amount of Earthlight

Two-Layer Model

- Use balance at boundary to space to find temperature of the glass
- Since energy in from Sun and energy out from glass are same as bare rock model, temperature is too!

$F_{in} = F_{out}$

Boundary $\frac{S(1-\alpha)}{4} = \sigma T_{glass}^4$

Glass $\left(\frac{S(1-\alpha)}{4\sigma}\right)^{1/4} = T_{glass}$

$255\text{ K} = T_{glass}$

Two-Layer Model

- Recall from boundary balance that $S(1-\alpha)/4 = \sigma T_{glass}^4$
- Substitute this into surface balance
- Find that surface is **much warmer** under glass!
- $303\text{ K} = 30\text{ }^\circ\text{C} = 86\text{ }^\circ\text{F}$

$F_{in} = F_{out}$

Surface $\frac{S(1-\alpha)}{4} + \sigma T_{glass}^4 = \sigma T_{sfc}^4$

$\phi T_{glass}^4 + \phi T_{glass}^4 = \phi T_{sfc}^4$

$2T_{glass}^4 = T_{sfc}^4$

$\sqrt[4]{2} T_{glass} = T_{sfc} = 1.189 T_{glass} = 303\text{ K}$

Summary

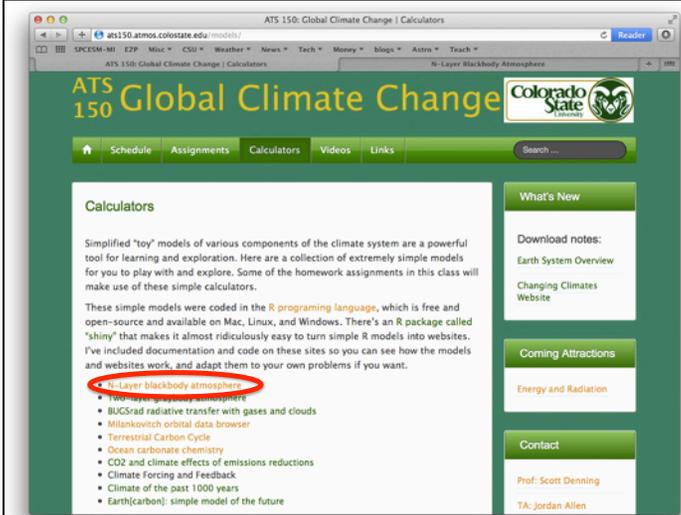
- Layer models are not meant to be accurate or predictive, just to help us understand how the world works
- Energy balance of “**bare rock**” model is **way too cold** to support life!
- Adding a layer of **glass** makes the surface nice and **toasty**
- **Where did the heat come from?**

Bathtub Analogy

- If faucet runs faster than drain, level rises
- And vice versa
- Drain runs faster when water is deep
- Adding glass to layer model acts like a clog in the drain
- Water rises until drainage = inflow again



Faucet ~ Sun
Drain ~ Thermal emission
Water level ~ temperature

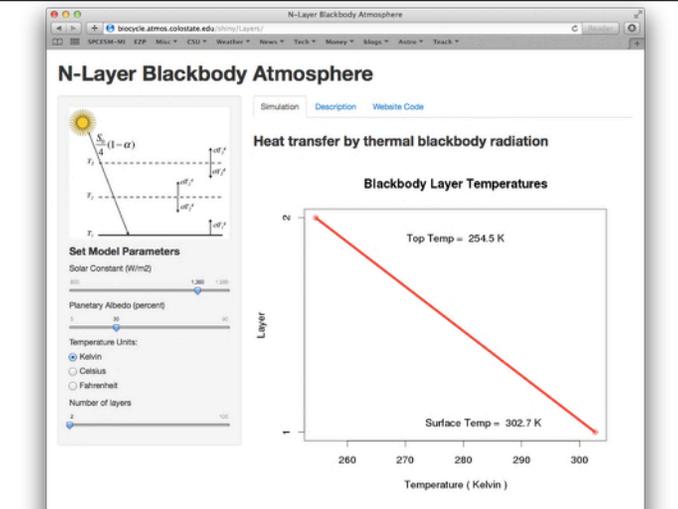


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N-Layer blackbody atmosphere

- BUGSrad radiative transfer with gases and clouds
- Milankovitch orbital data browser
- Terrestrial Carbon Cycle
- Ocean carbonate chemistry
- CO2 and climate effects of emissions reductions
- Climate Forcing and Feedback
- Climate of the past 1000 years
- Earth(carbon): simple model of the future



N-Layer Blackbody Atmosphere

Heat transfer by thermal blackbody radiation

Blackbody Layer Temperatures

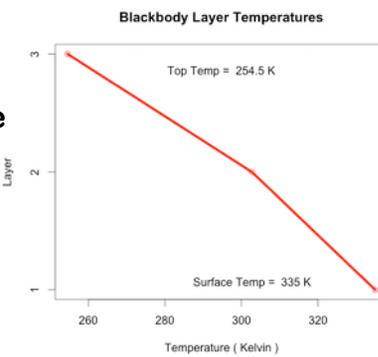
Top Temp = 254.5 K

Surface Temp = 302.7 K

Temperature (Kelvin)

Three Layers

- Top boundary (“skin”) temperature is always the same
- As we add layers, the surface gets hotter
- (bathtub drain gets slower)



Blackbody Layer Temperatures

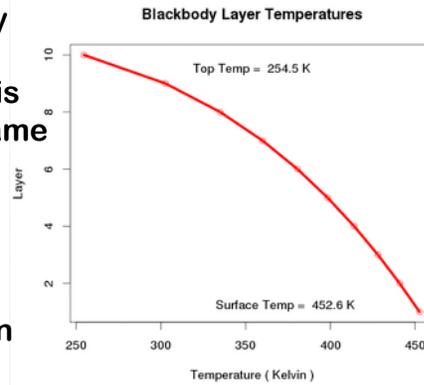
Top Temp = 254.5 K

Surface Temp = 335 K

Temperature (Kelvin)

Ten Layers

- Top boundary (“skin”) temperature is always the same
- As we add layers, the surface gets hotter
- (bathtub drain gets slower)



100 Layers!

- Top boundary (“skin”) temperature is always the same
- As we add layers, the surface gets hotter
- (bathtub drain gets slower)

