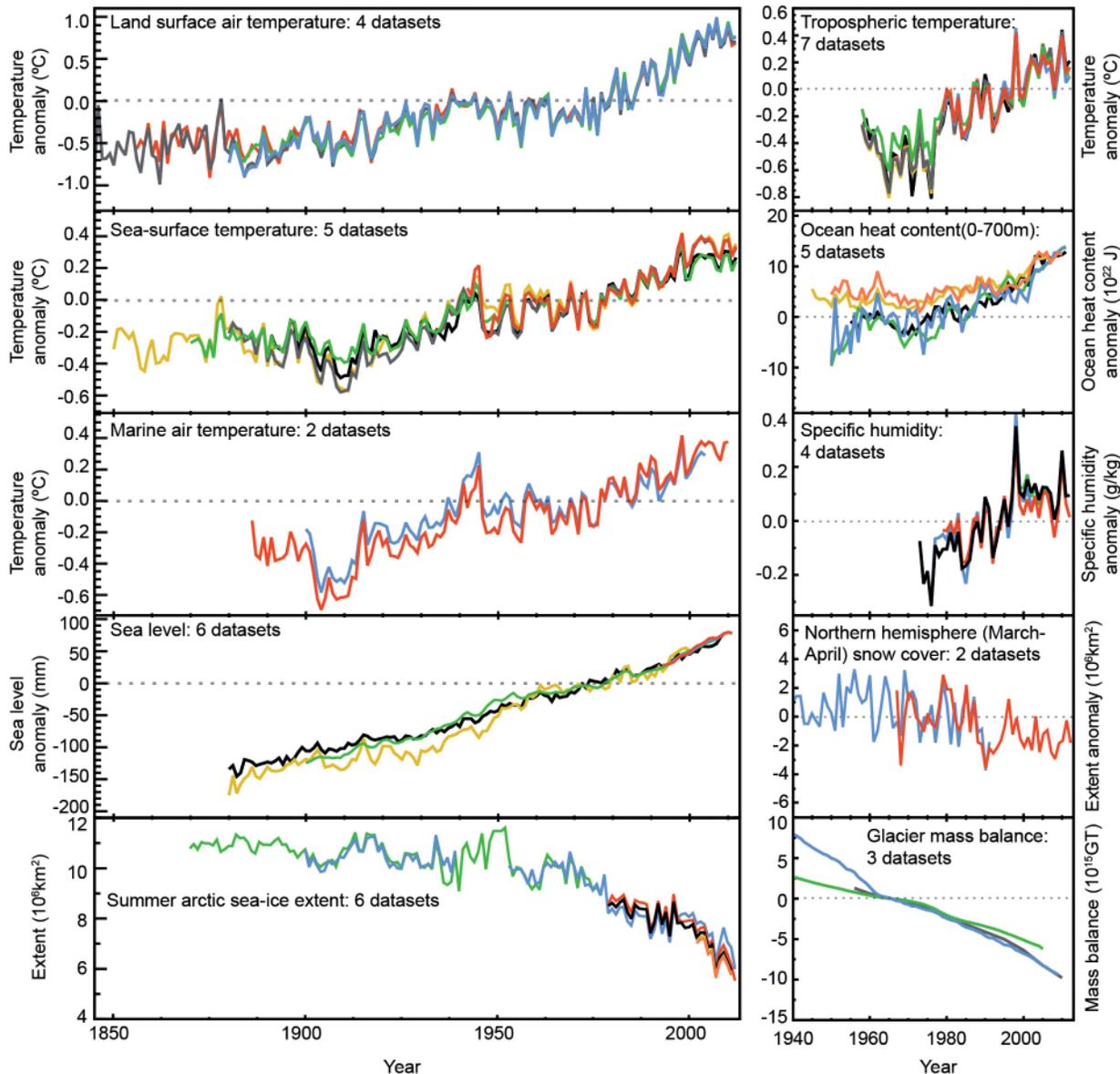


# Course Outline

## Climate 201: Modern Climate Change

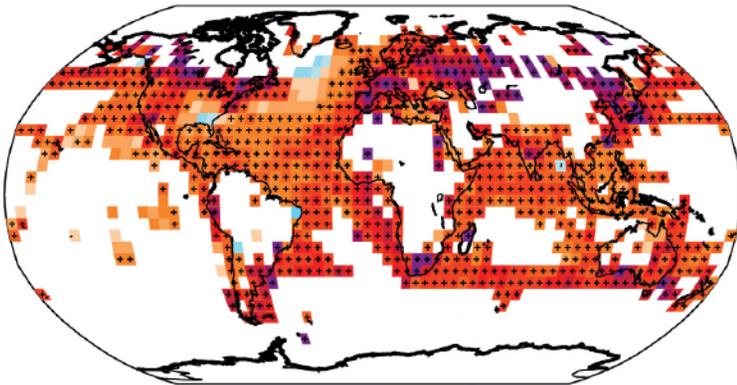
- 2/4 Climate Change in a Nutshell
- 2/11 Forcing, Feedback, & Sensitivity
- **2/18 What to Expect in the Future**
- 2/25 Impacts of Climate Change
- 3/4 Mitigation, Adaptation, & Costs

# Historical Trends

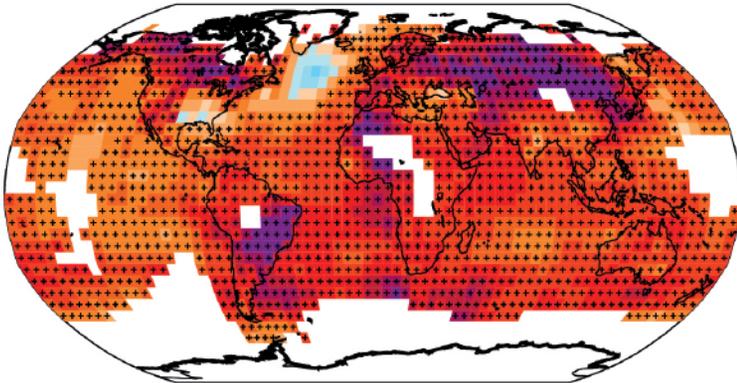


- Observed changes are large and observed in many different ways
- Warmer in air, on surface, and in ocean
- Less ice both on land and floating on sea
- Rising seas, moister air

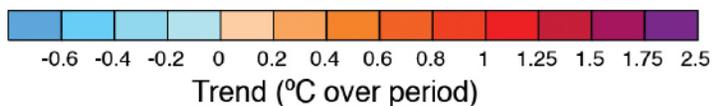
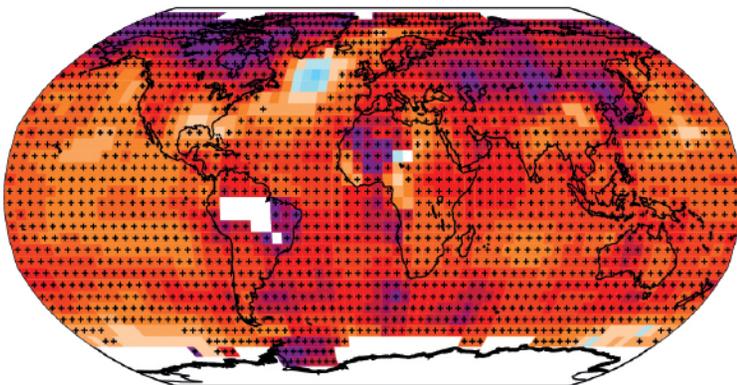
HadCRUT4 1901-2012



MLOST 1901-2012



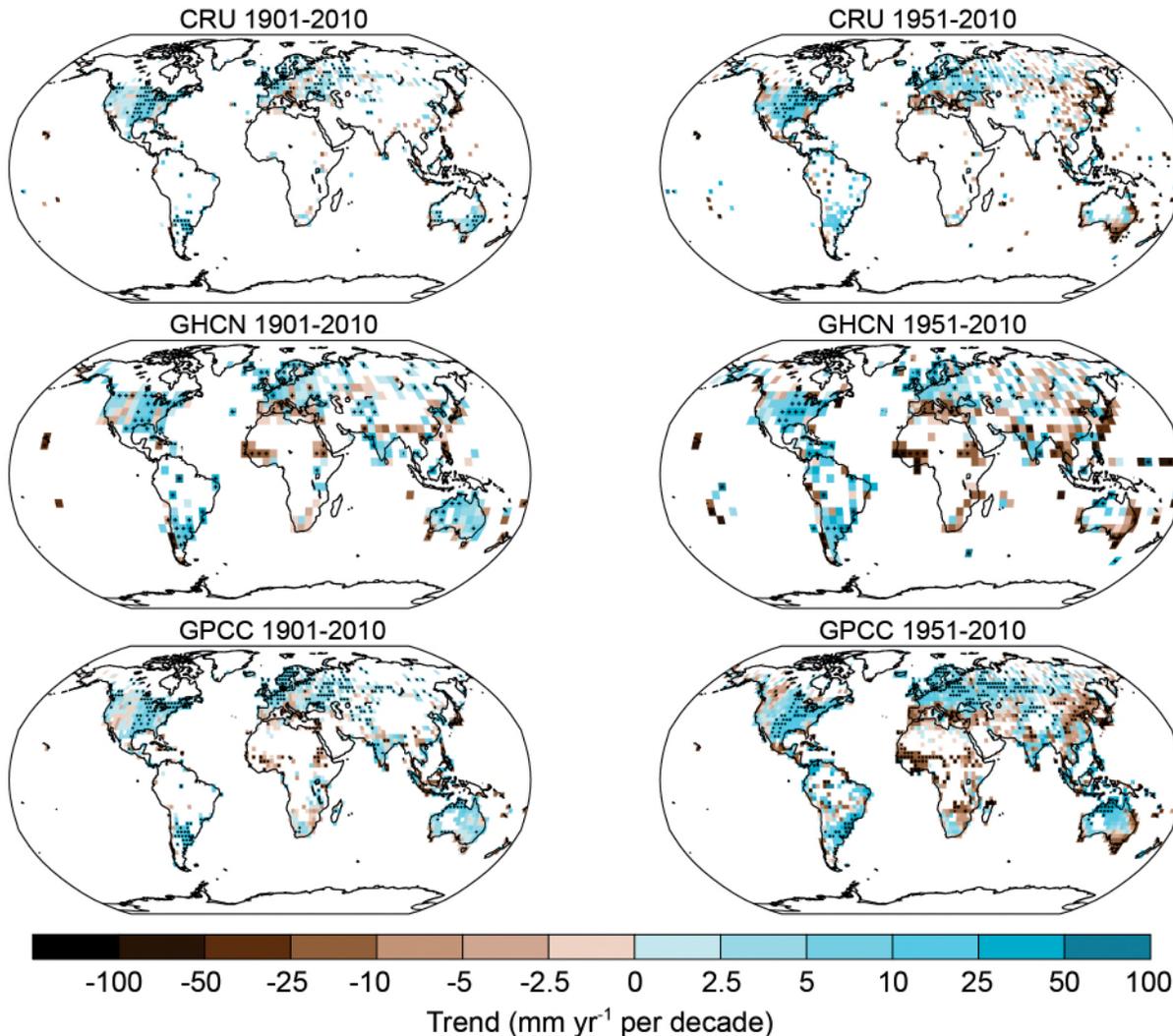
GISS 1901-2012



# Historical Trends

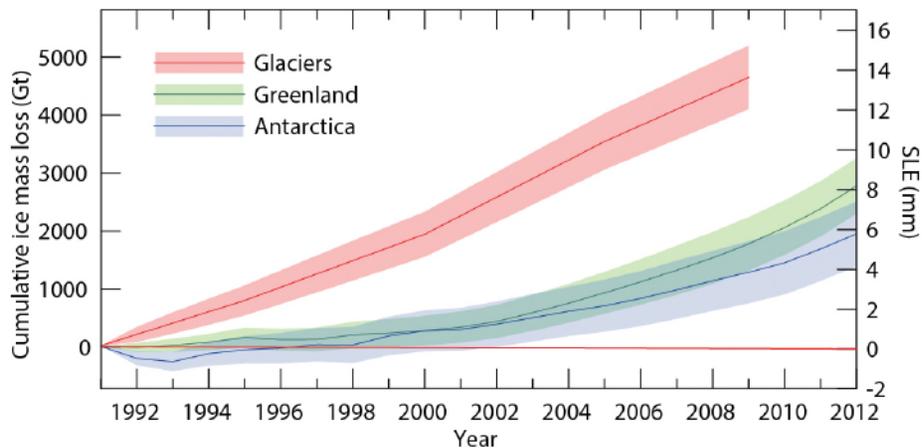
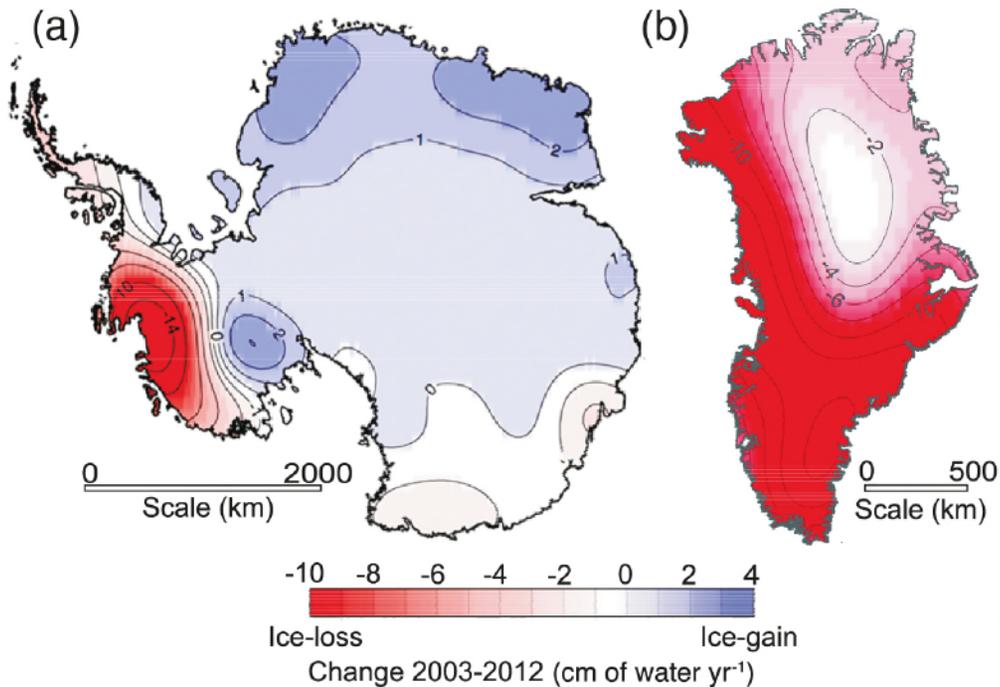
- CO<sub>2</sub> increased by about 33% (295 ppm in 1900, 397 ppm in 2012)
- Warming almost everywhere
- Most warming
  - on land
  - In NH
  - In Arctic
- Global average warming of about 0.8 °C (1.3 °F)
- Land warming ~ 50% more

# Historical Precipitation Trends



- Warmer air evaporates more water
- Overall precipitation must therefore also increase
- Wet places get wetter, and dry places get drier

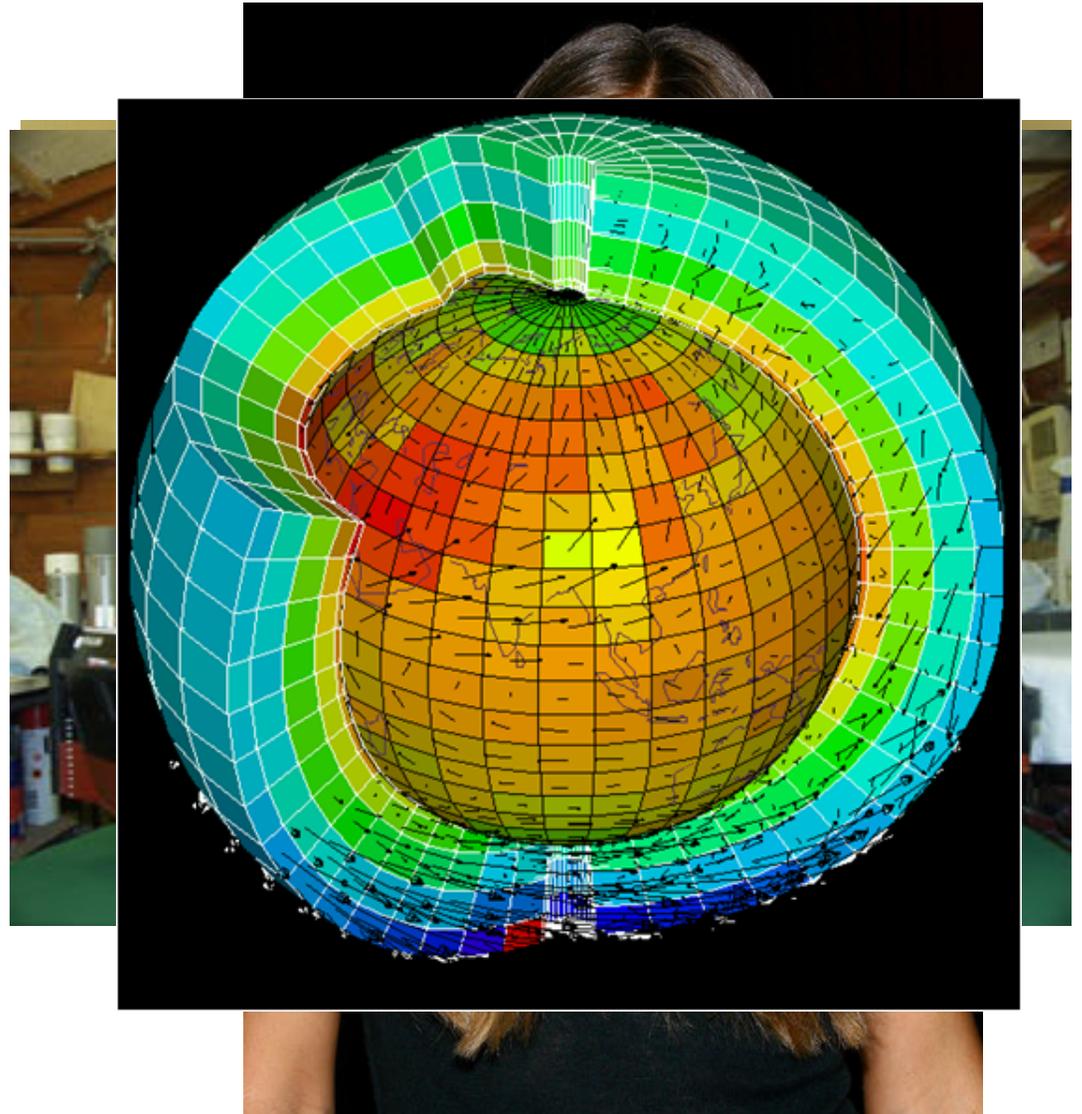
# Ice Mass Loss (observed)



- Good data only since the 1990s (sat)
- Smaller glaciers are losing mass faster than ice sheets
- Greenland and West Antarctic Ice Sheets are losing mass
- East Antarctic Ice Sheet (much bigger) is gaining mass

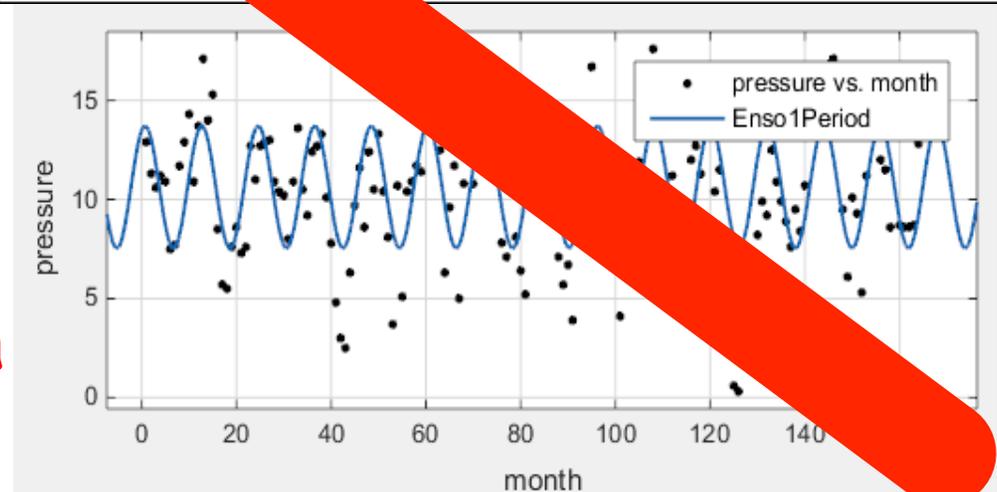
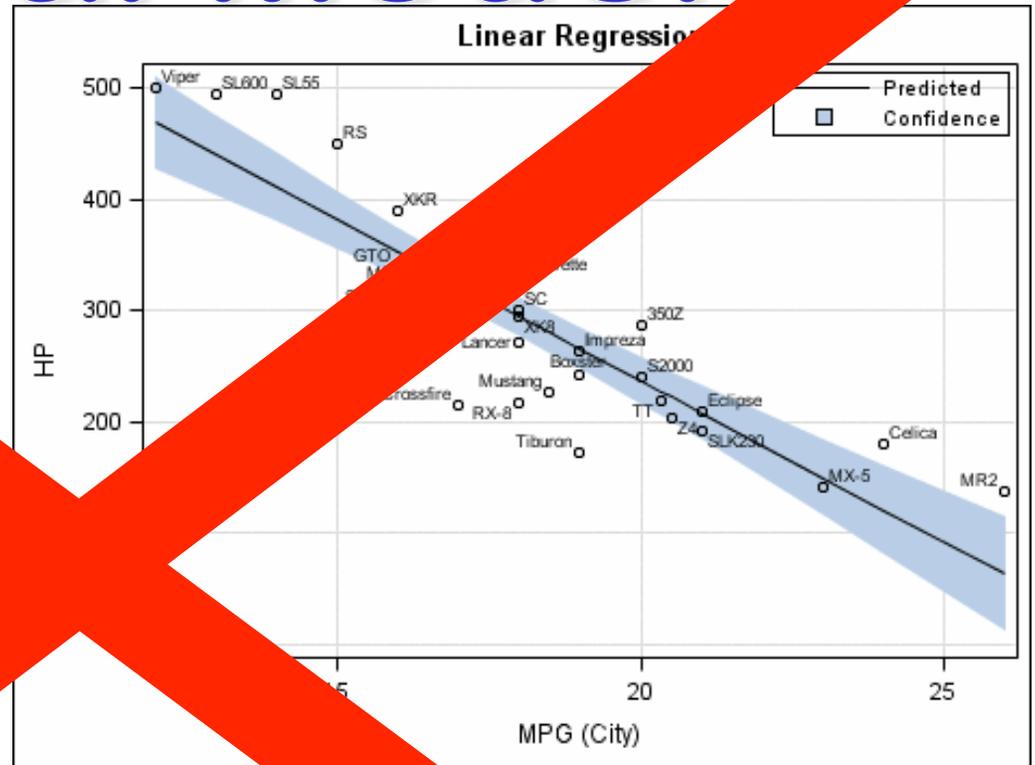
# Climate Models

- What is a “model”
- What does it mean to model the climate?
- How do modern climate models work?
- How good are they?
- What can they tell us?
- What can't they tell us?

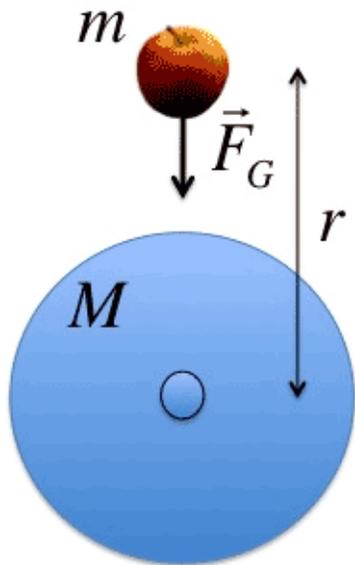


# Empirical Models

- Generalized mathematical formulation with adjustable coefficients
- Combinations of
  - polynomials
  - exponential growth & decay
  - Periodic waves and cosine
- Coefficients fit to data (e.g., least squares)



# Deterministic Models



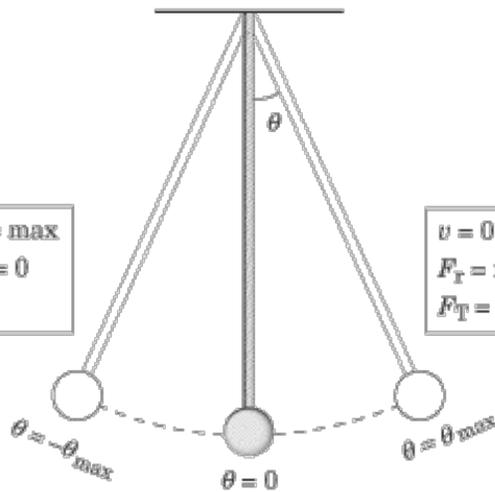
$$F = ma$$

$$\frac{GMm}{r^2} = ma$$

Then, cancelling  $m$  on both sides:

$$a = \frac{GM}{r^2} = g$$

$$\begin{array}{ll} v = 0 & U_g = \max \\ F_r = \max & KE = 0 \\ F_T = \min & \end{array}$$



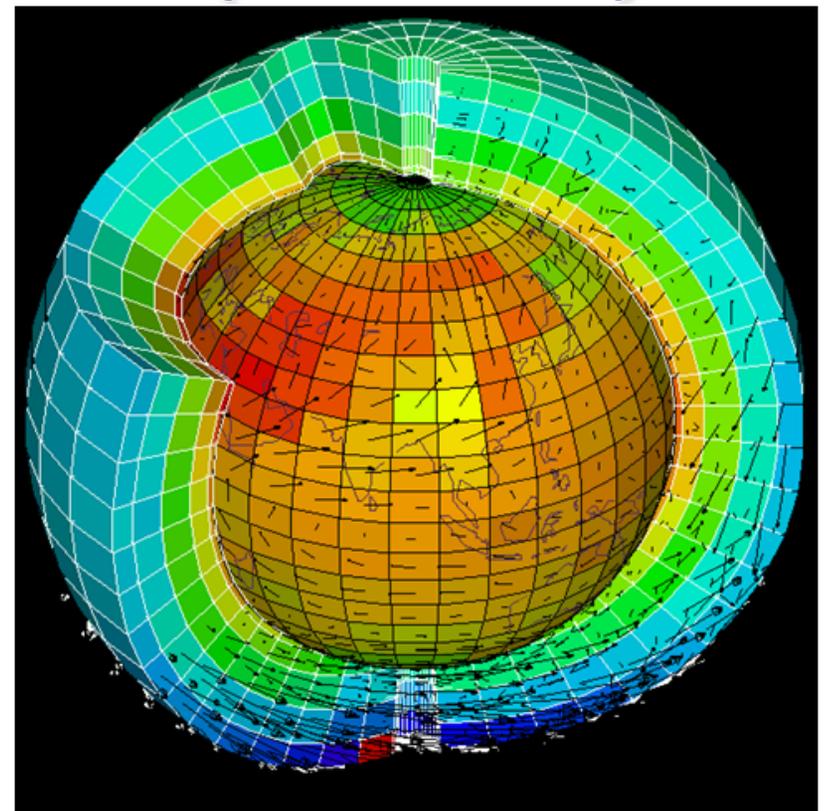
$$\begin{array}{ll} v = 0 & U_g = \max \\ F_r = \max & KE = 0 \\ F_T = \min & \end{array}$$

$$\begin{array}{ll} v = \max & U_g = \min \\ F_r = 0 & KE = \max \\ F_T = \max & \end{array}$$

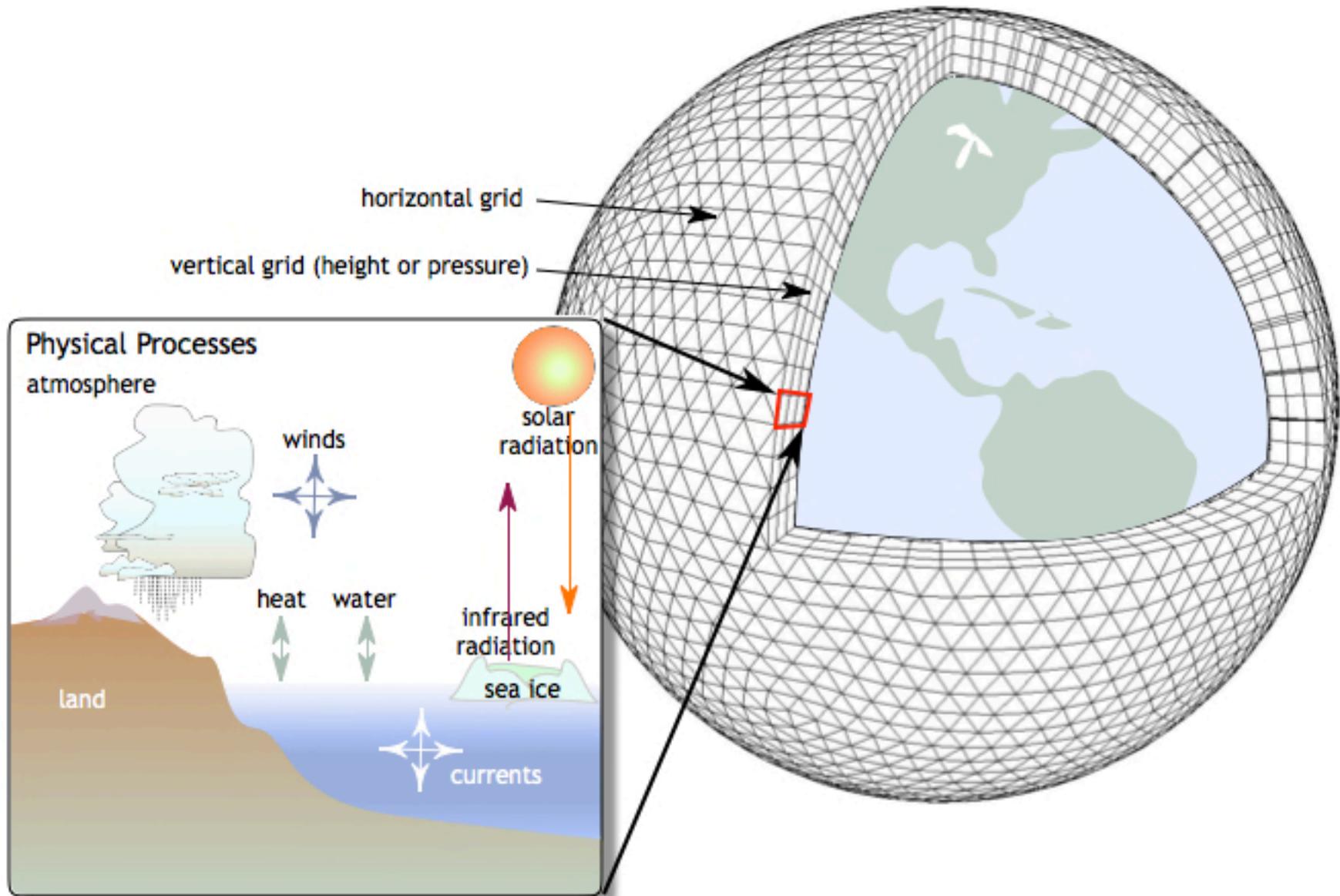
- Formulated as **“cause and effect”**
- Common in physics and chemistry
- Usually take the form of **differential equations**
- **Initial & boundary-value** problems
- May still have adjustable coefficients

# “General Circulation Models” (GCMs)

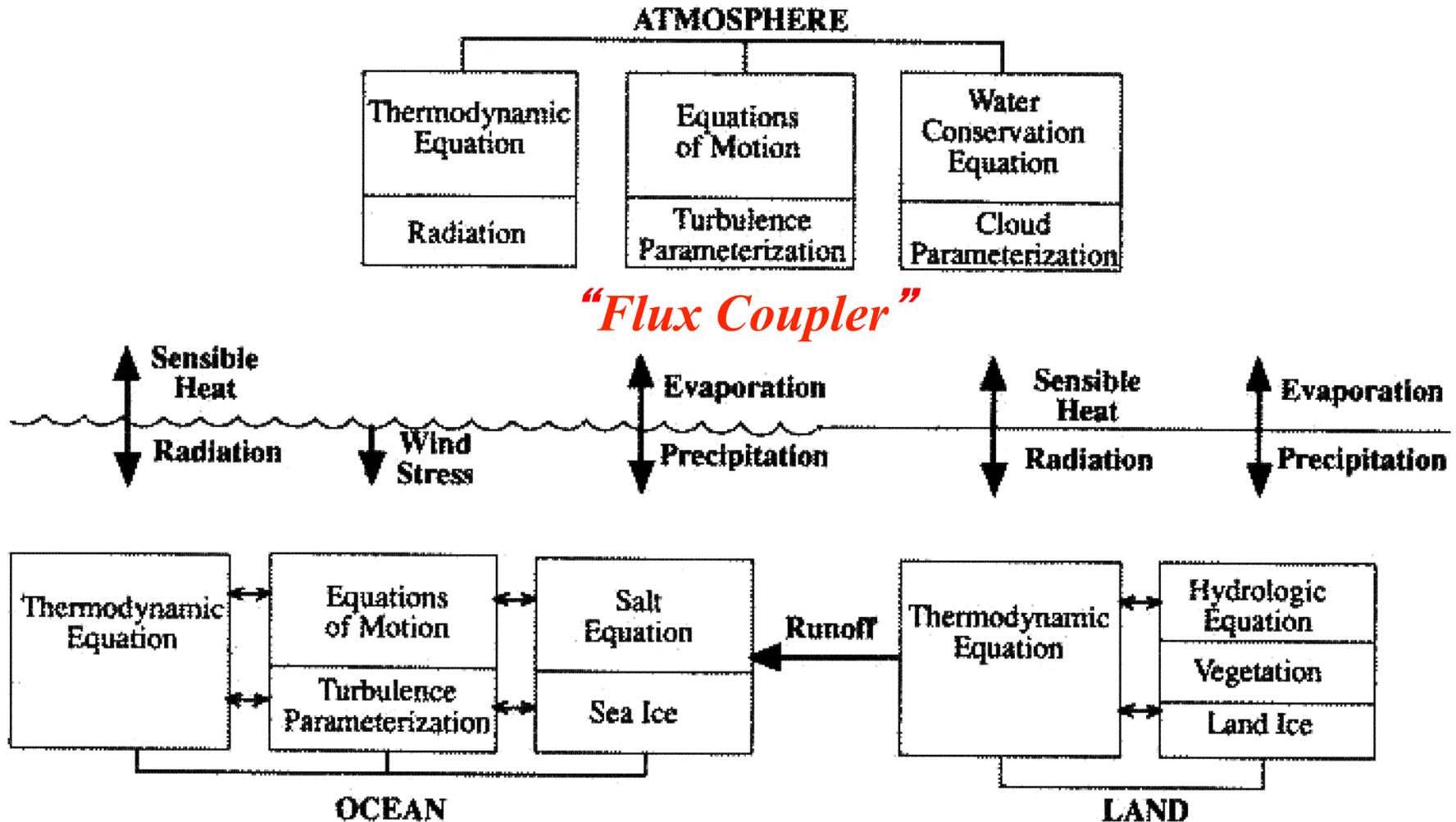
- **Deterministic, not empirical**
- “ $F = ma$  of a compressible fluid on a rotating sphere with radiation, thermodynamics, and phase transitions”
- Allow detailed prediction of future states at high resolution in both space and time
- **Same equations:**
  - Weather forecasting (initial value problem)
  - Climate simulation (boundary value problem)



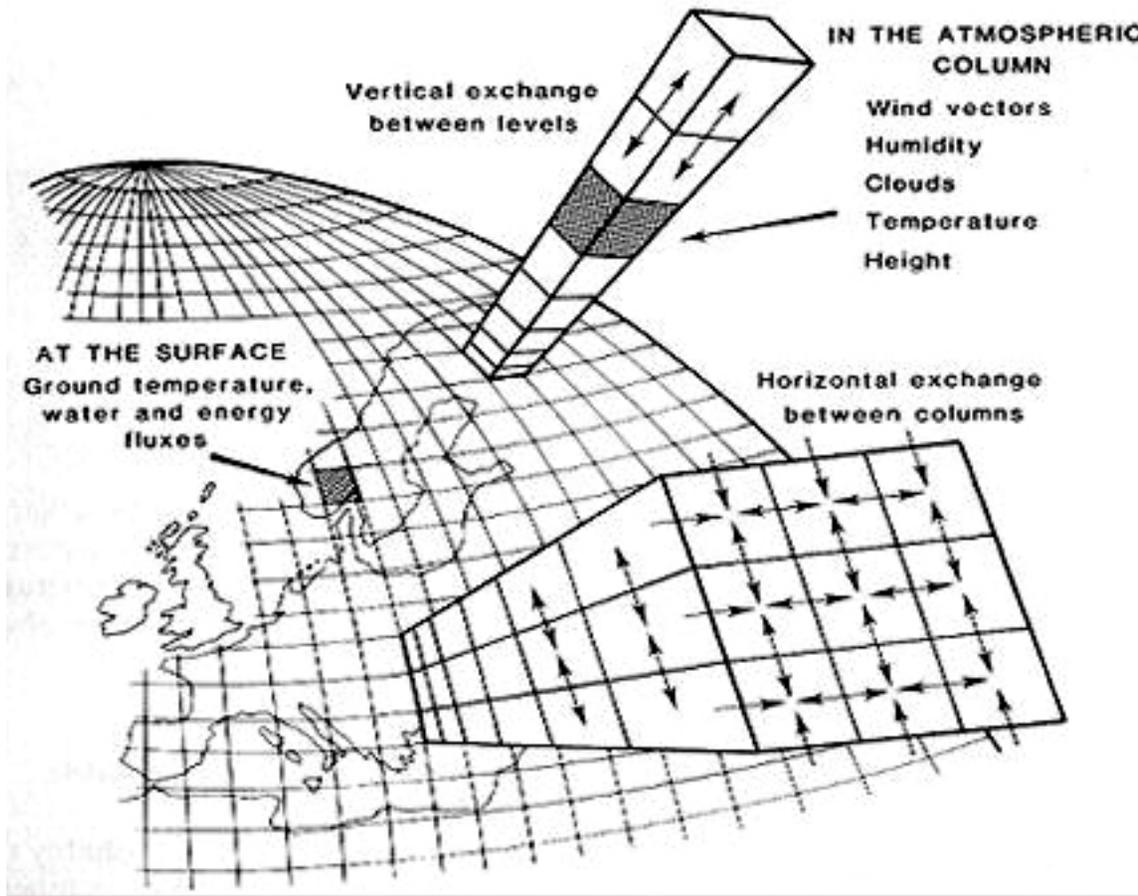
# Climate Model Processes



# Climate Model Structure



# Climate Model Grids

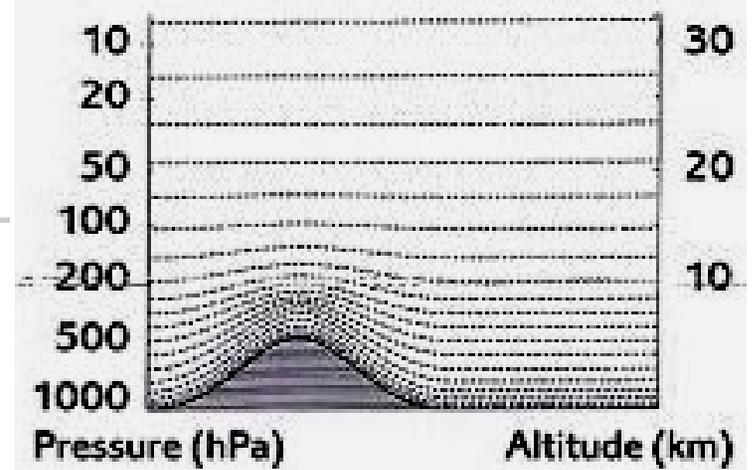


Typical climate model

$\Delta x \sim 100$  km

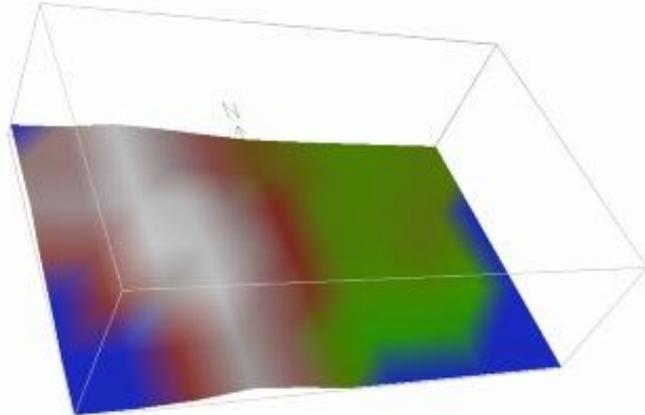
Typical weather forecast model

$\Delta x \sim 12$  km



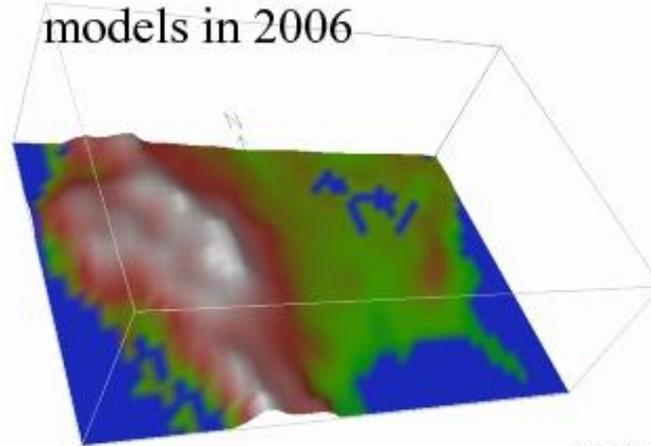
# Topography at Different Resolutions

Climate Models circa early 1990s



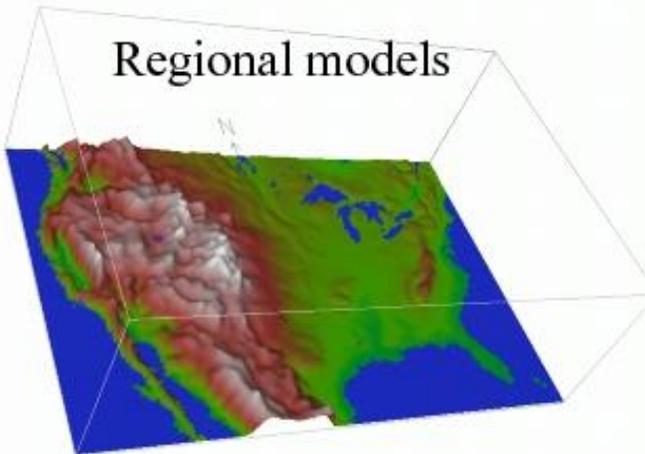
400 km

Global coupled climate models in 2006



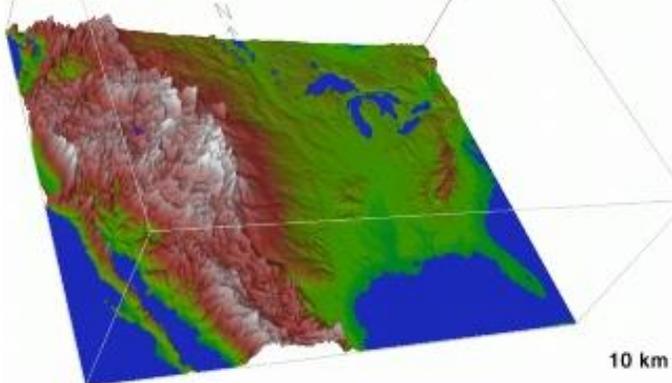
100 km

Regional models



25 km

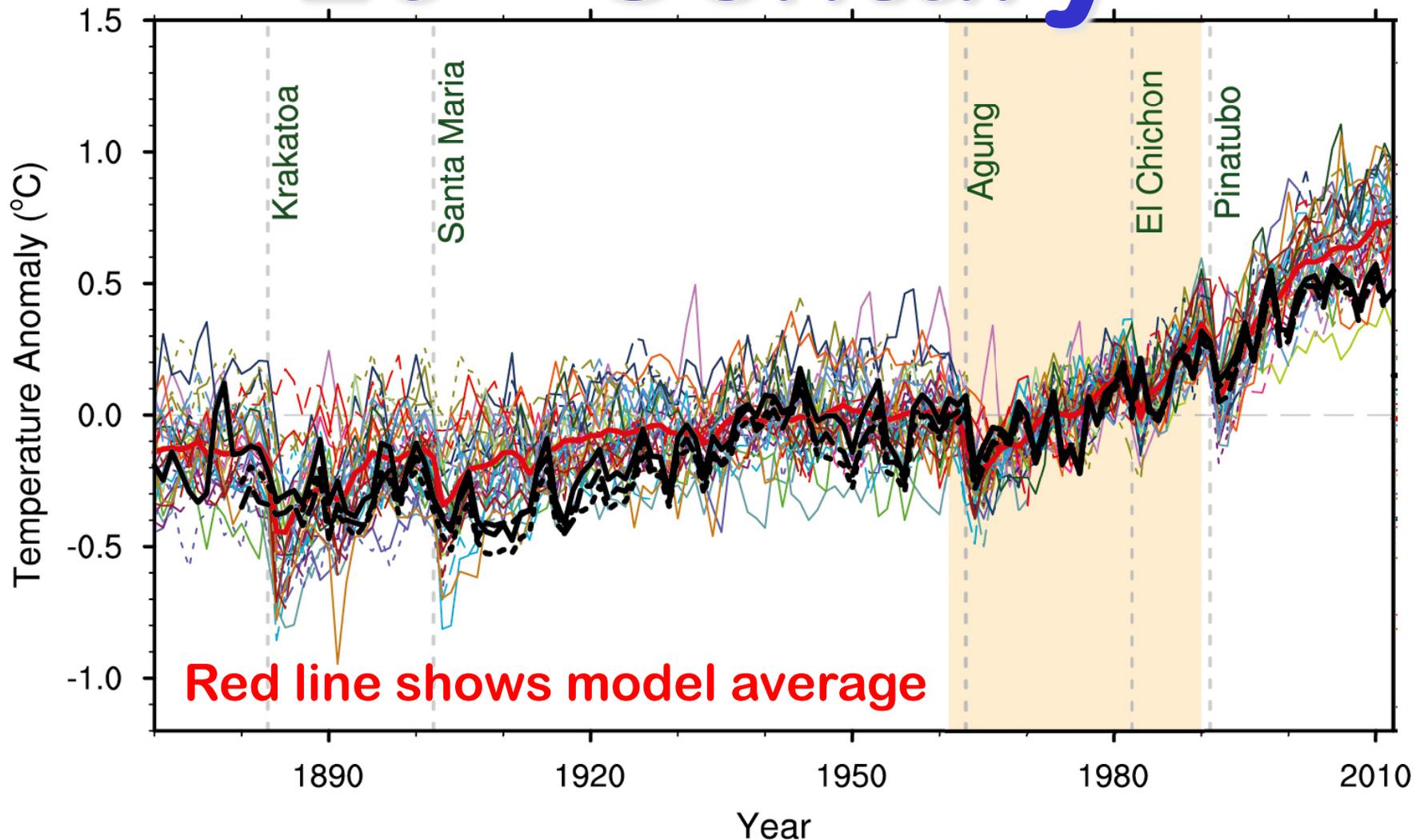
Global models in 5-10 yrs



10 km

Optimistic view on model-development

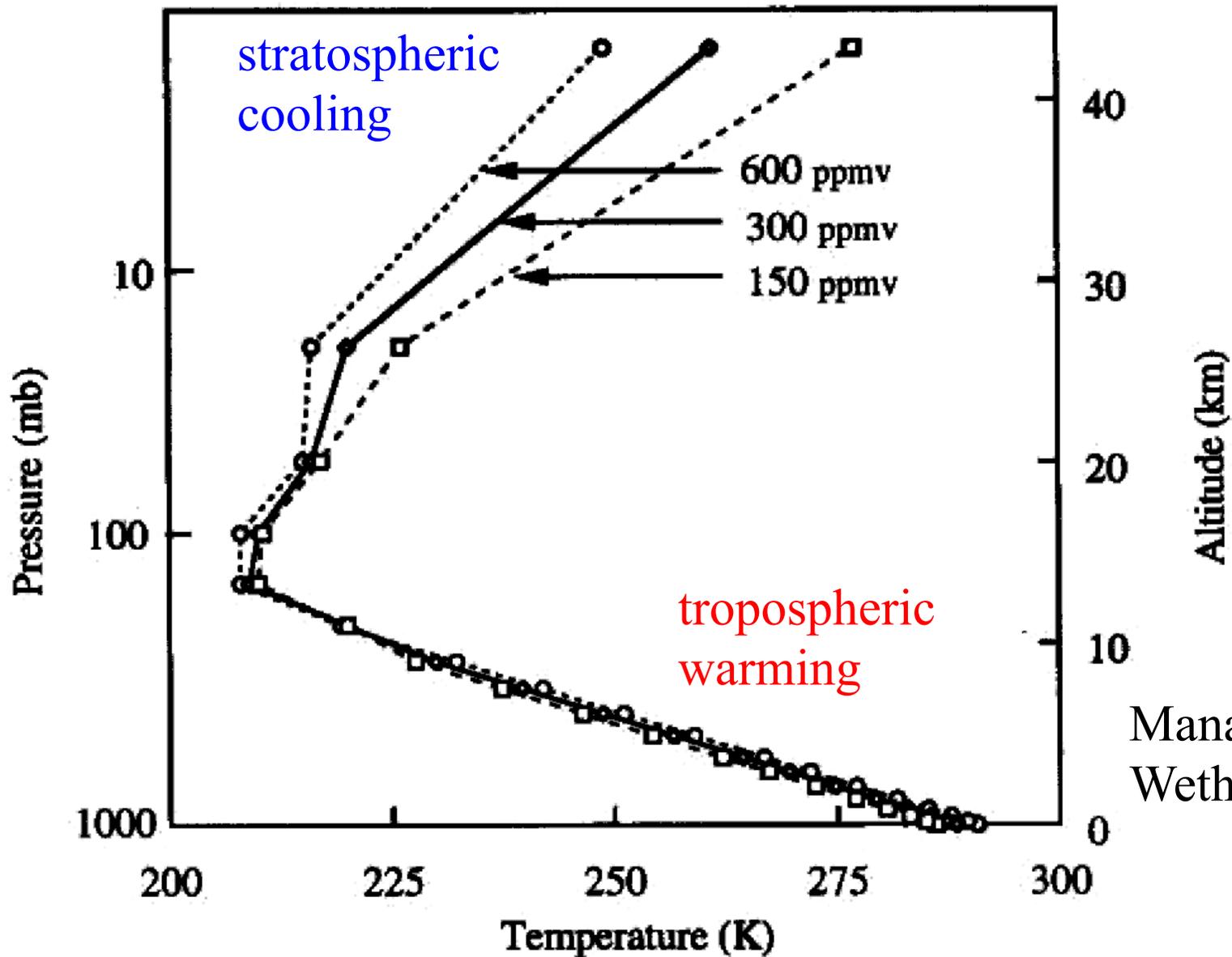
# 20<sup>th</sup> Century



**3 observation sets  
36 GCMs**

**Observations almost identical  
Lots of model variations**

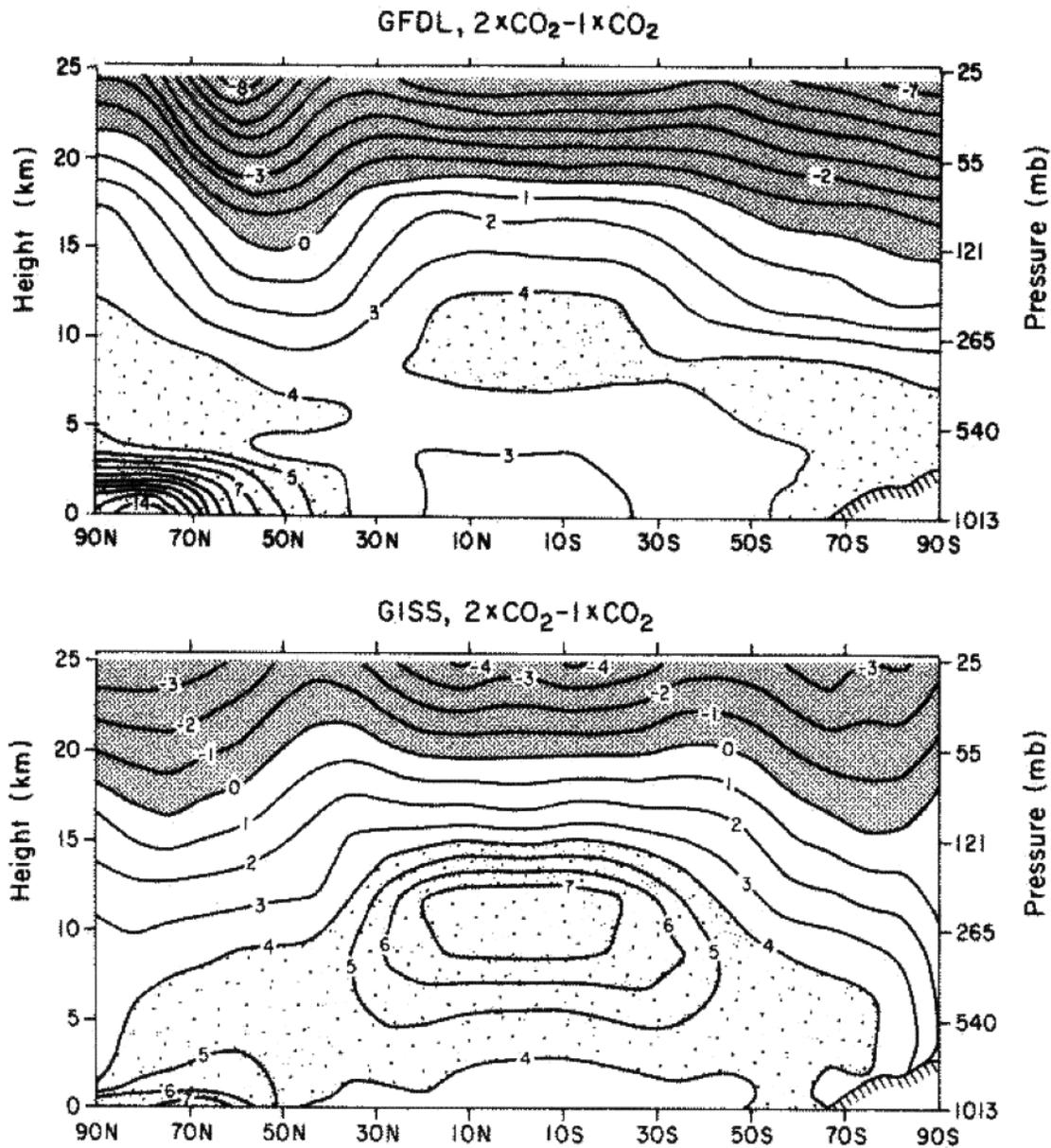
# Warming under Cooling



Manabe and Wetherald (1967)

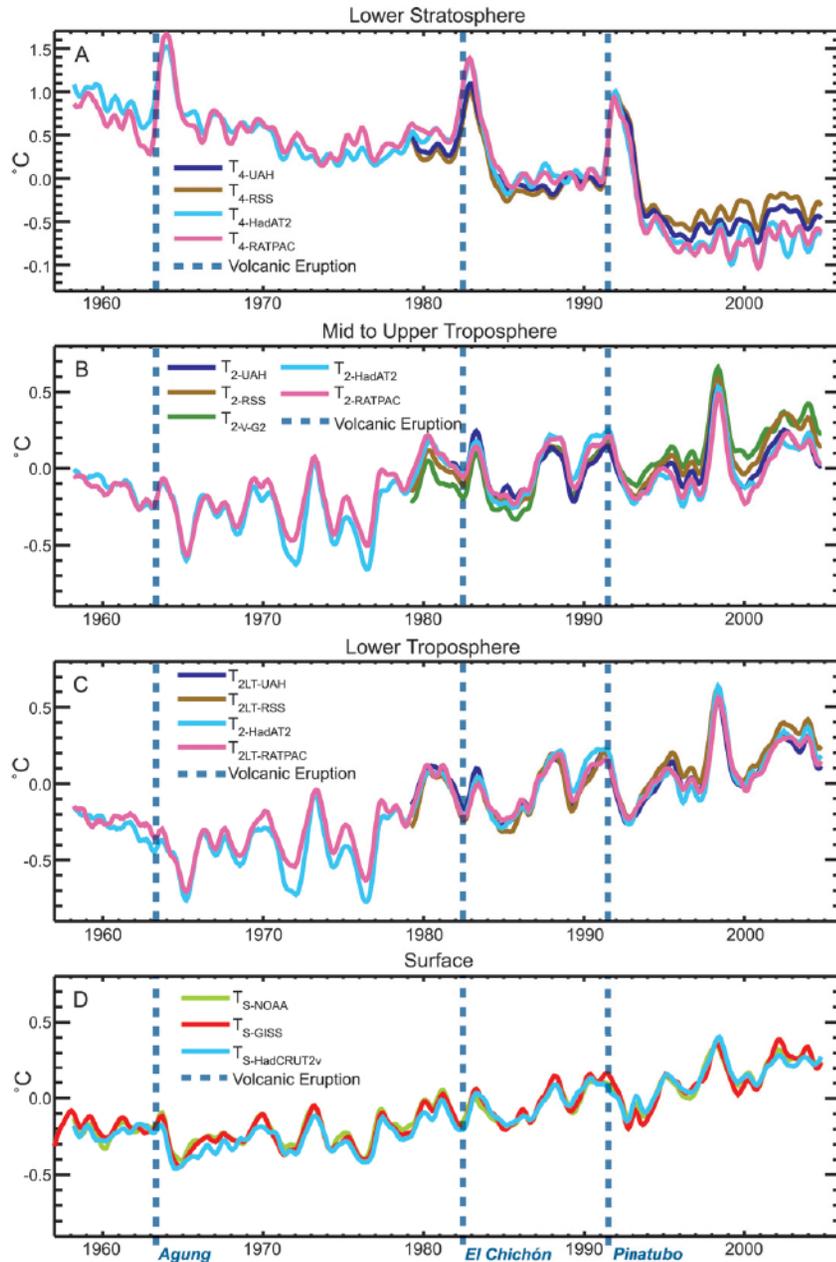
# Predicted Vertical Structure

- Greenhouse “signature” is tropospheric warming and stratospheric cooling
- Predicted in mid-1980’s by climate models

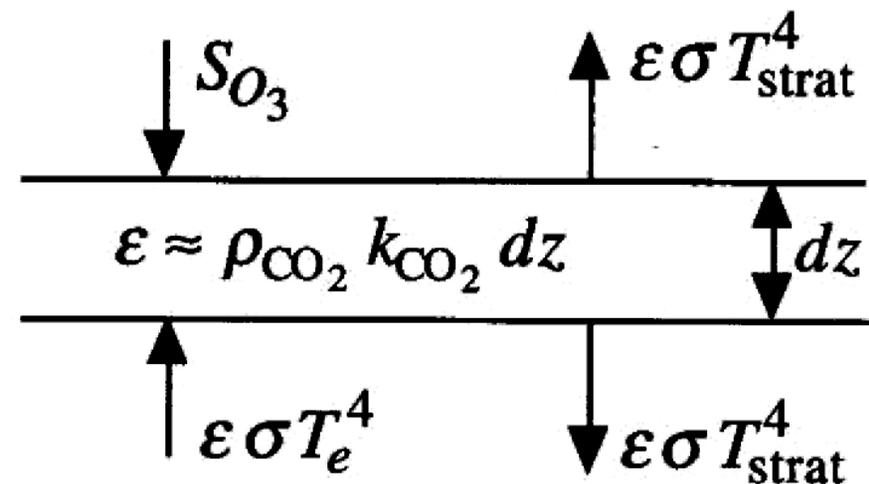


**Fig. 12.9** Contour plot of the zonally averaged change in air temperature during DJF resulting from a CO<sub>2</sub> doubling in two models that each give a global-average surface temperature increase of 4°C. Cooling and warming greater than 4°C are shaded. [Top panel, Wetherald and Manabe (1986), reprinted with permission from Kluwer Academic Publishers; bottom panel, Hansen *et al.* (1984), © American Geophysical Union, as printed in Schlesinger and Mitchell (1987), © American Geophysical Union.]

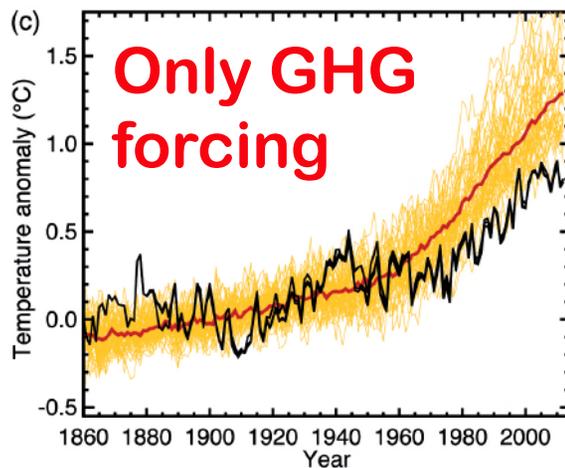
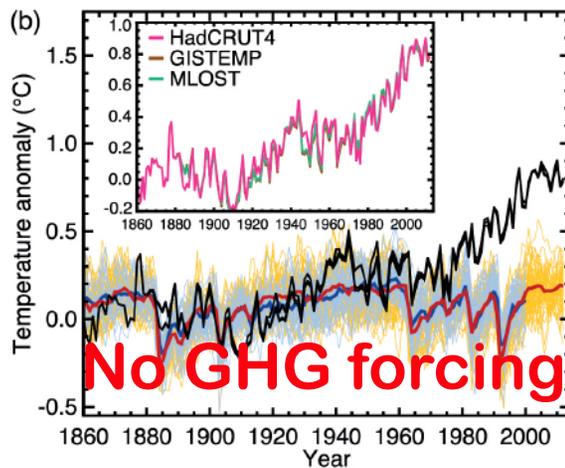
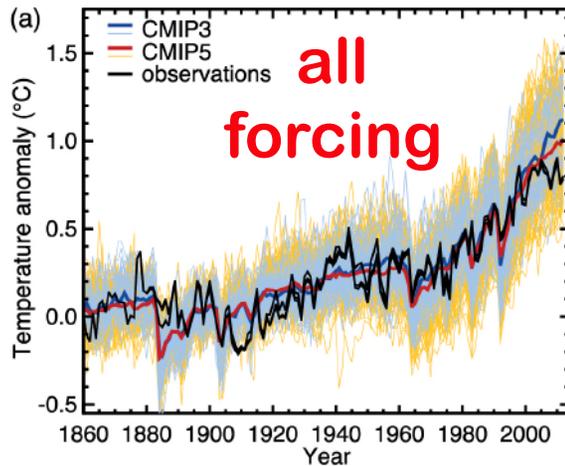
## OBSERVED AIR TEMPERATURES



# Balloons, Satellites, and Surface Obs



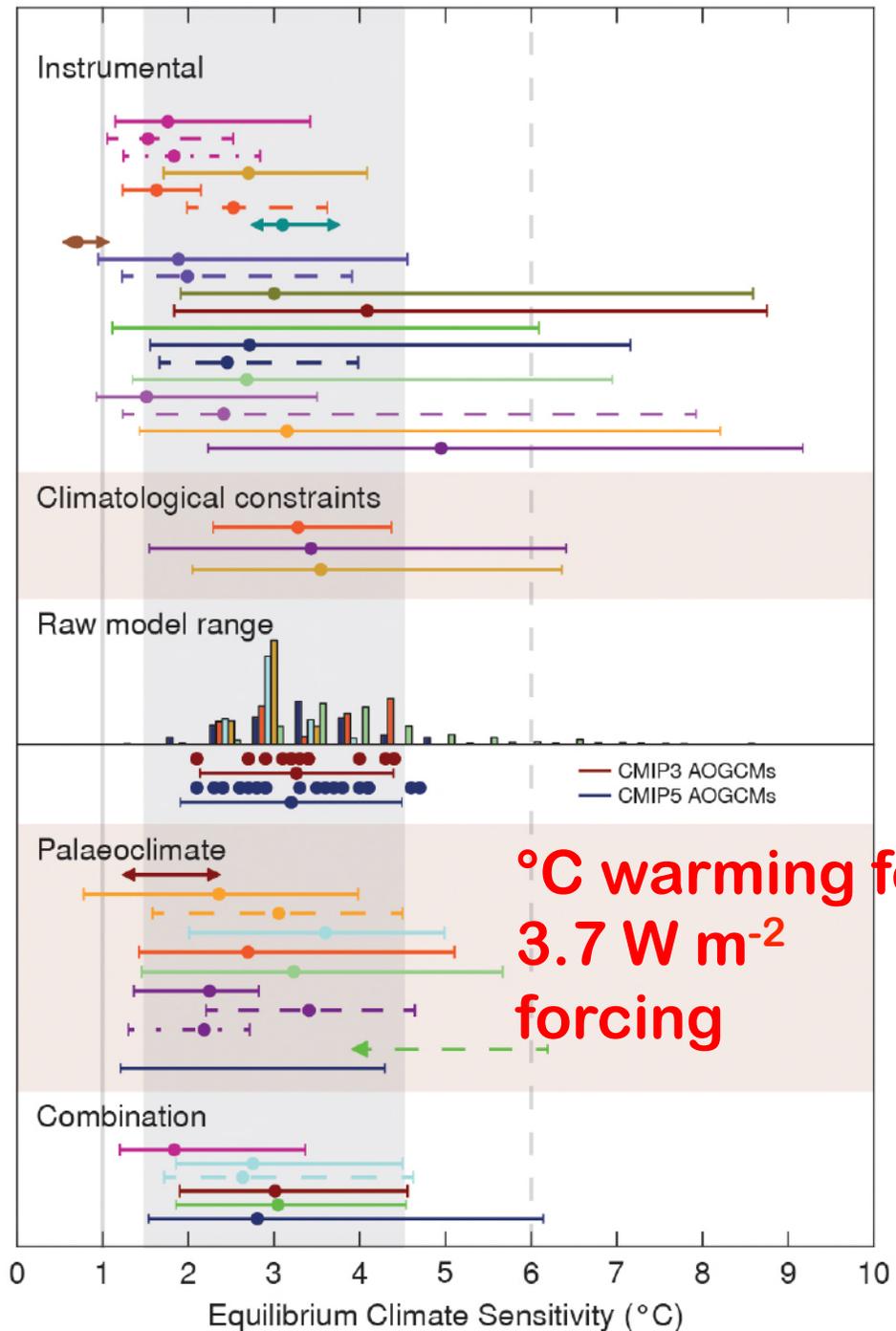
# Hindcasts of 20<sup>th</sup> Century



- Models without greenhouse forcing don't predict enough warming
- Models with only greenhouse forcing predict too much warming
- Models with all forcing do a good job of predicting past climate change

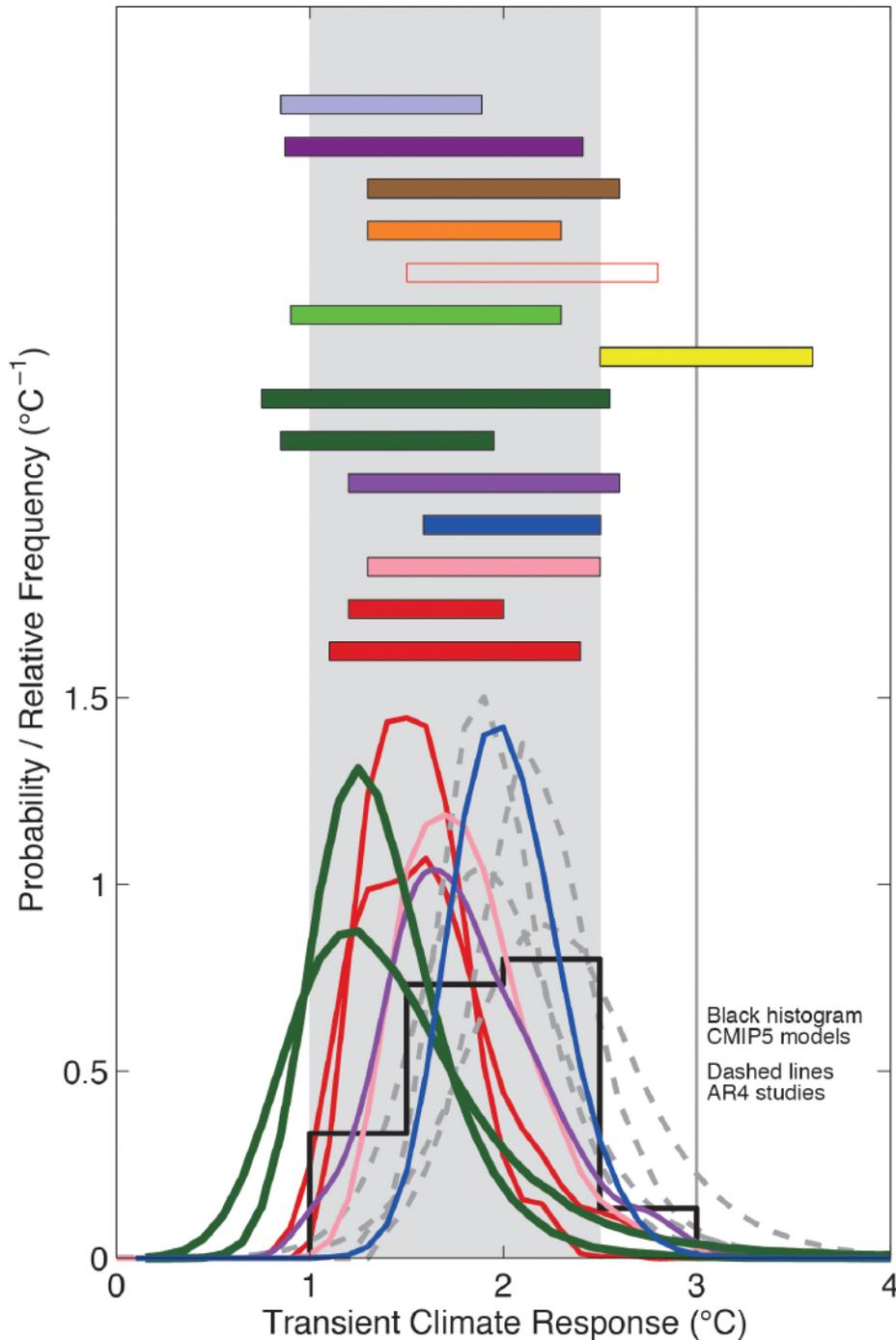
# Equilibrium Climate Sensitivity

- Surface warming required to re-establish thermal **equilibrium** at top of atmosphere
- Many lines of paleoclimate evidence and most GCMs find about **3 °C**



# Transient Climate Response

- Warming takes a long time because much of the heat is absorbed by the oceans
- TCR is **warming at time when CO<sub>2</sub> reaches 560 ppm**
- Models and obs show **TCR ~ 1 to 2.5 °C**

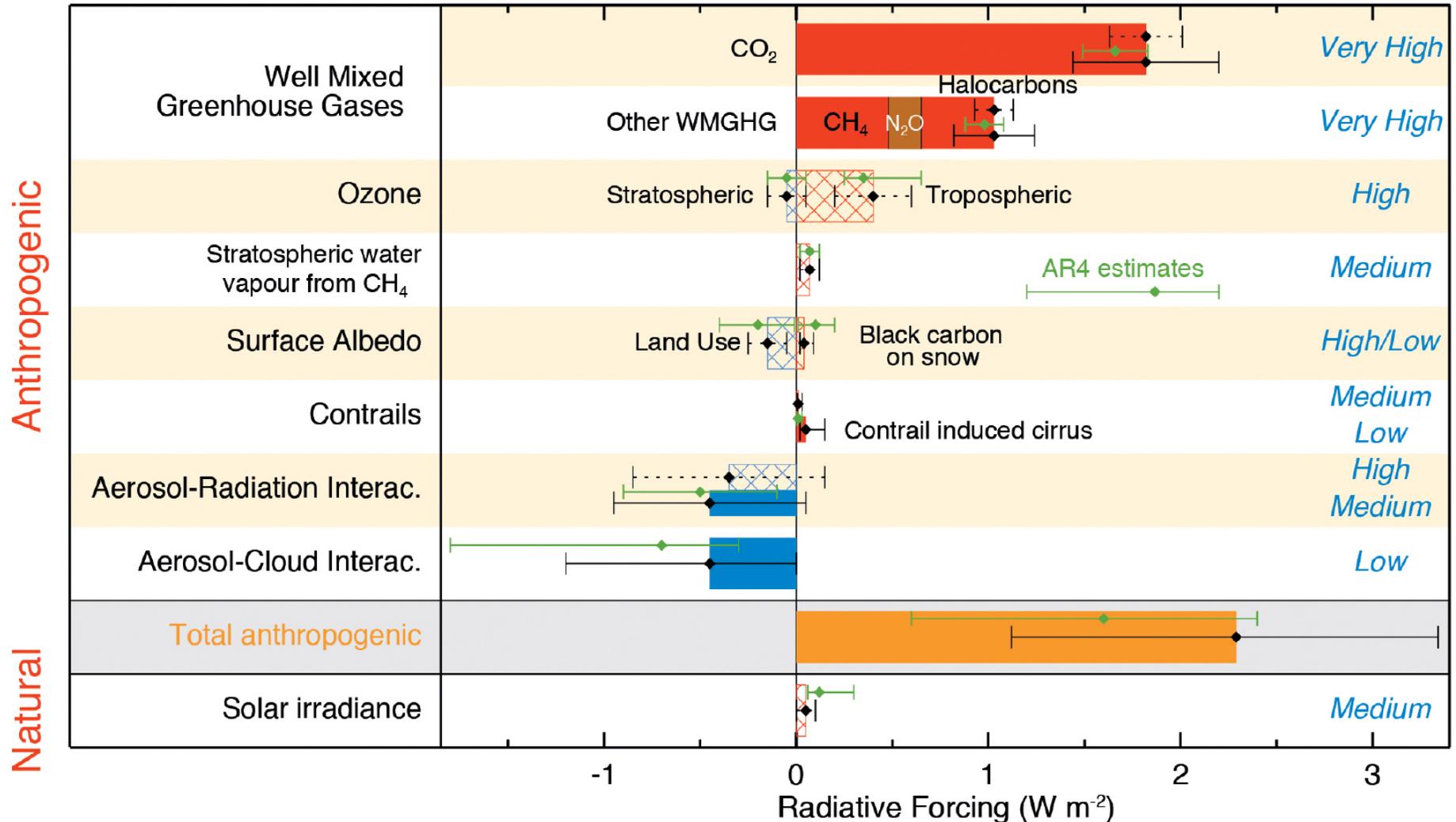


# Modern Climate Forcing

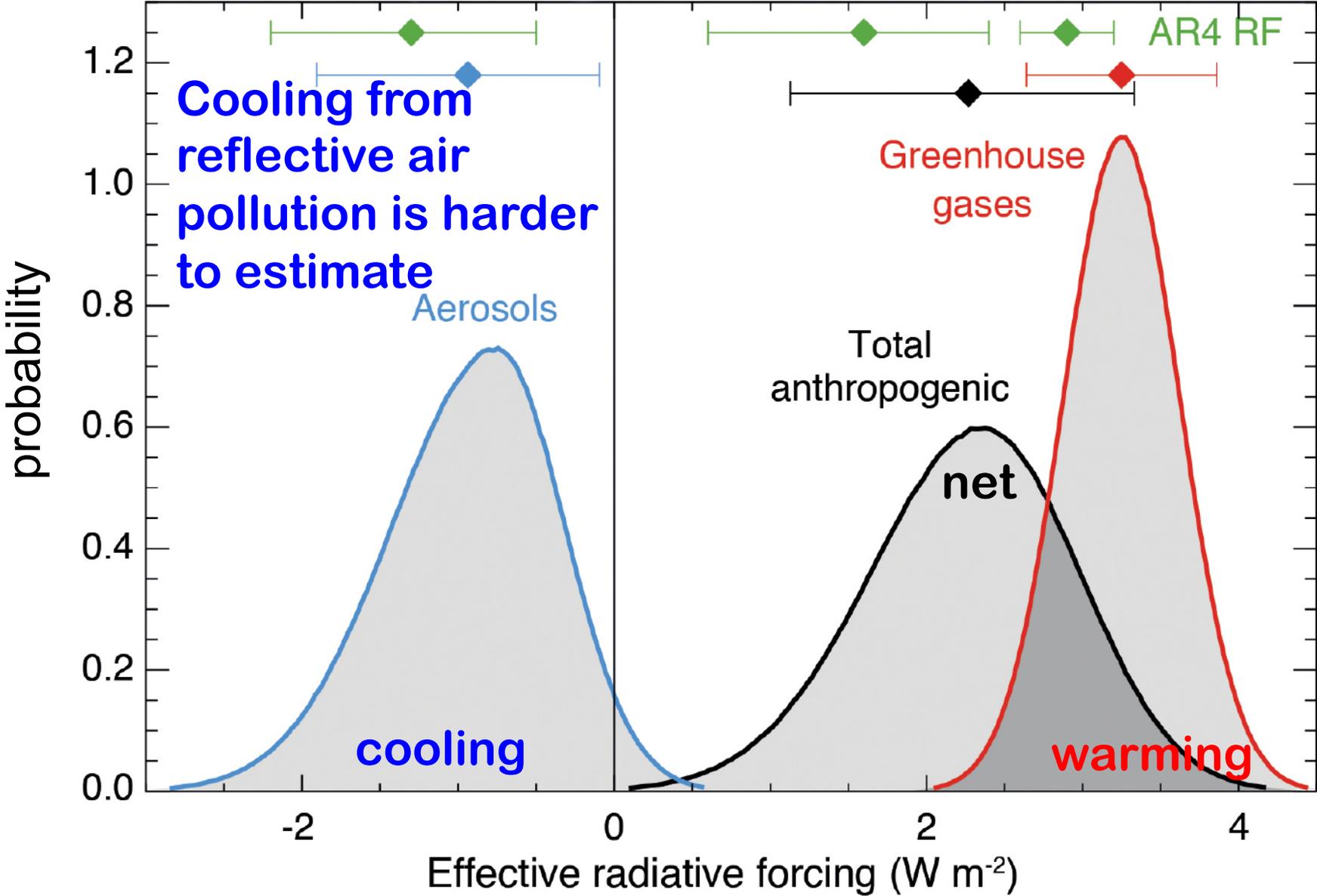
Radiative forcing of climate between 1750 and 2011

Forcing agent

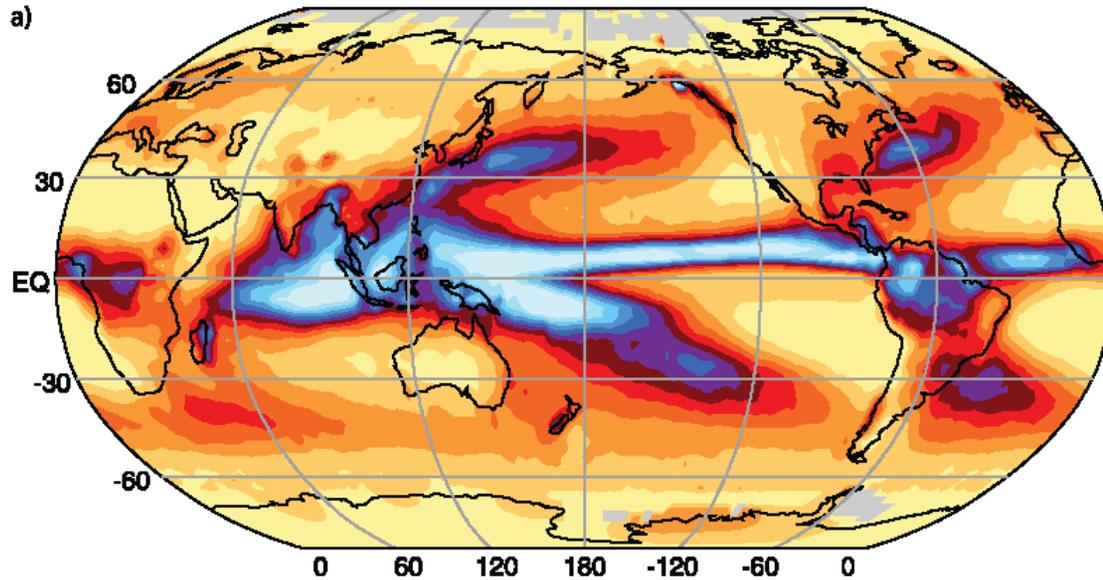
Confidence Level



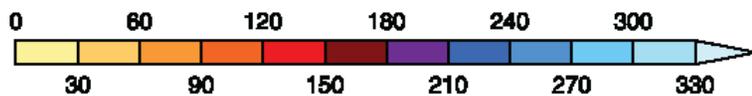
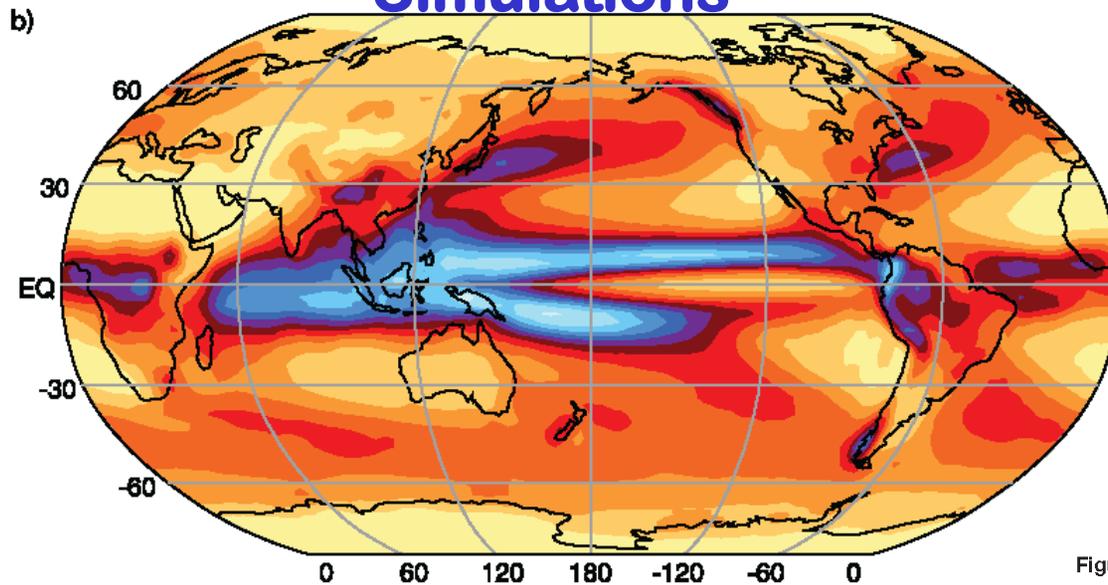
# Anthropogenic Forcing



## Observations



## Simulations



(cm)

# Model Evaluation: Precipitation

- General patterns and magnitudes very well simulated
- Problems with mountain ranges and some finer-scale patterns

Figure 8.5. Annual mean precipitation (cm), observed (a) and simulated (b), based on the multi-model mean. The Climate Prediction Center Merged Analysis of Precipitation (CMAP; Xie and Arkin, 1997) observation-based climatology for 1980 to 1999 is shown, and the model results are for the same period in the 20th-century simulations in the MMD at PCMDI. In (a), observations were not available for the grey regions. Results for individual models can be seen in Supplementary Material, Figure S8.9.

# The “Kaya Identity”

CO<sub>2</sub>: CO<sub>2</sub> emissions resulting from human activities

E: Primary energy consumption

G: GDP

P: Population

Kaya Identity: Formula that represents the relationship between human activities and CO<sub>2</sub> emissions

$$CO_2 = \frac{CO_2}{E} \times \frac{E}{G} \times \frac{G}{P} \times P$$

CO<sub>2</sub> emissions per unit  
energy consumption

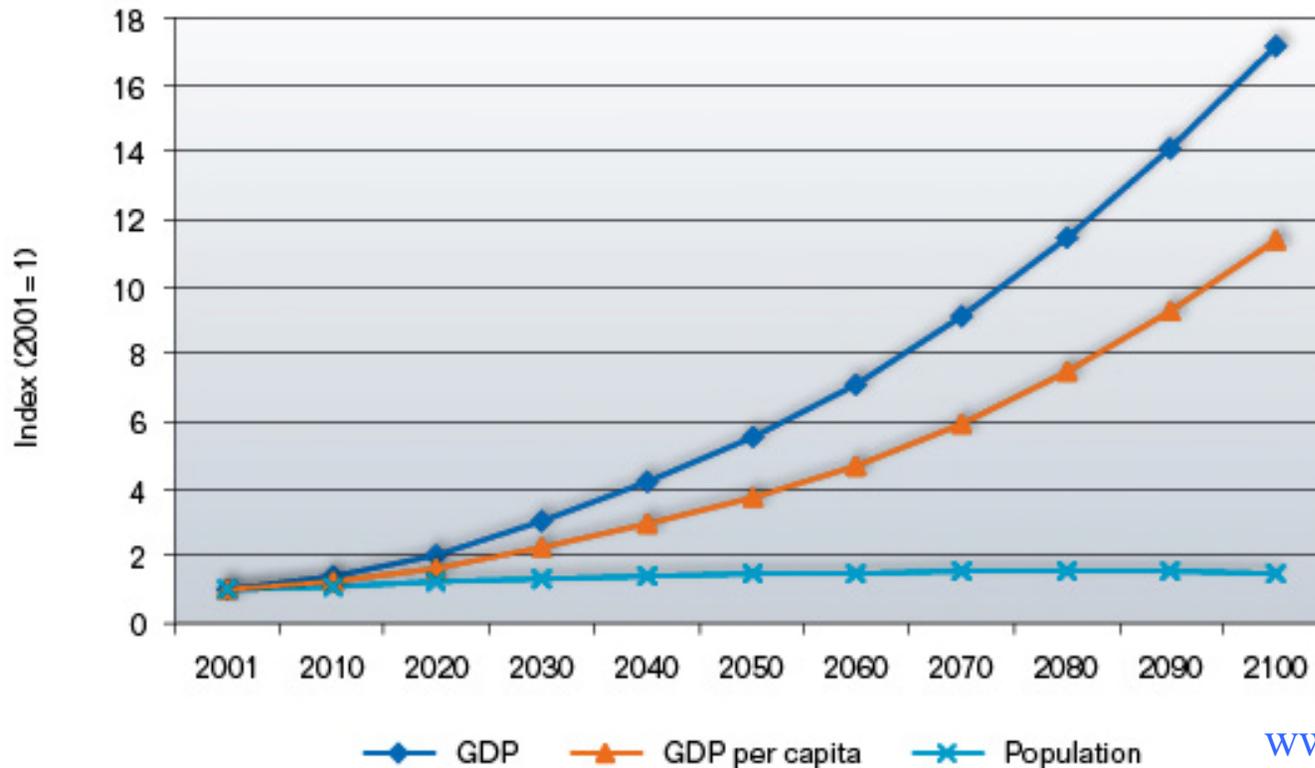
Energy efficiency of  
economic activities

Economic level  
per capita



- **Four factors determine future emissions:**
  - Population
  - Economic activity
  - Energy efficiency of economy
  - Carbon efficiency of energy

# Population is **not** the driver of future climate!

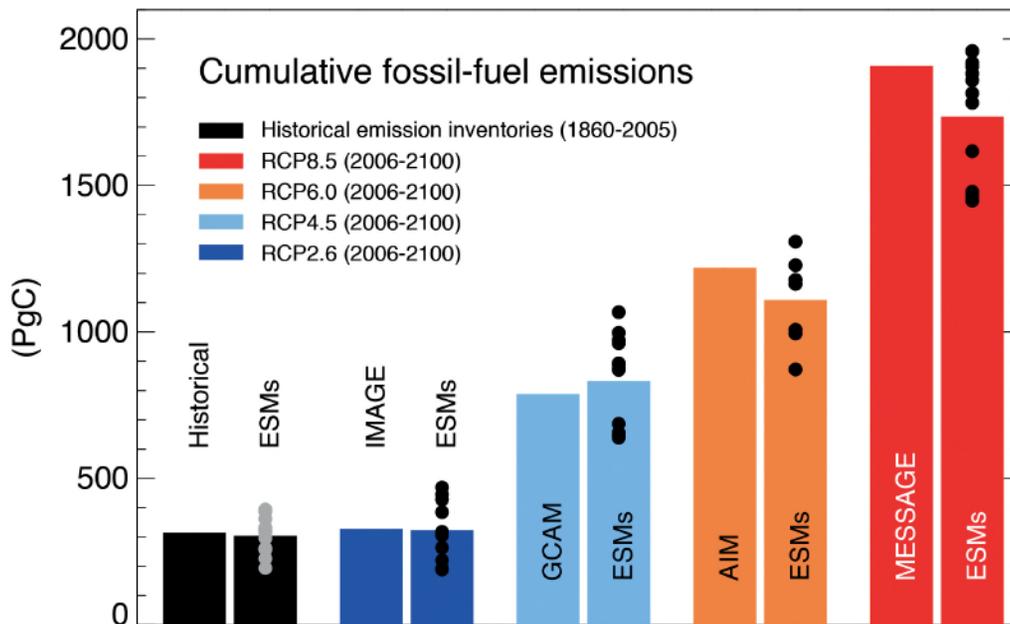
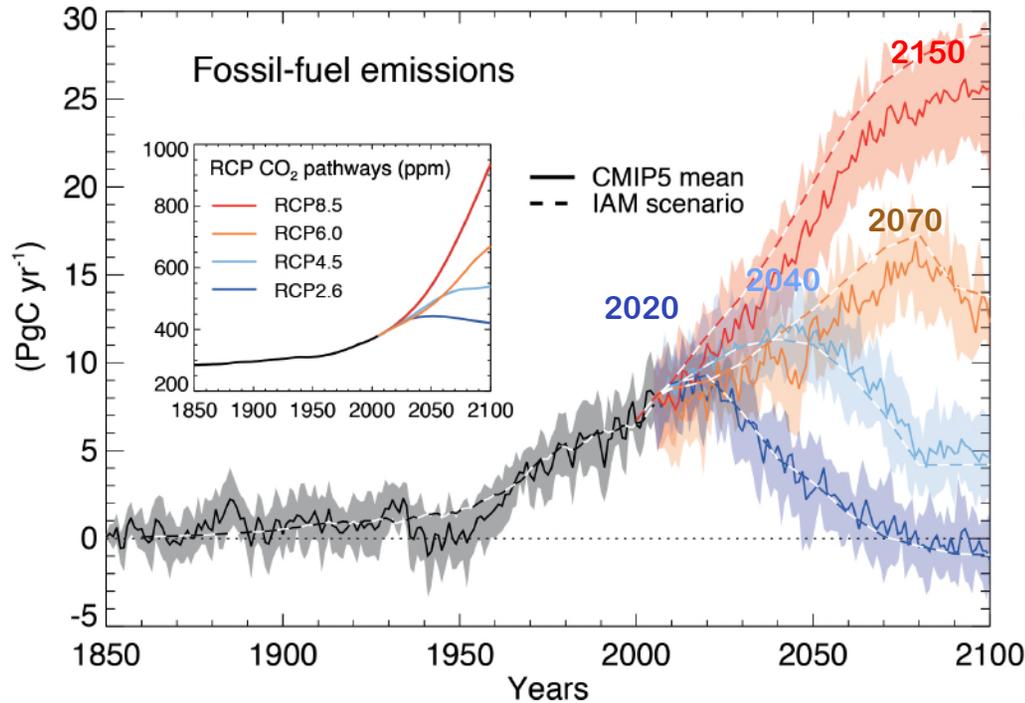


[www.garnautreview.org.au](http://www.garnautreview.org.au)

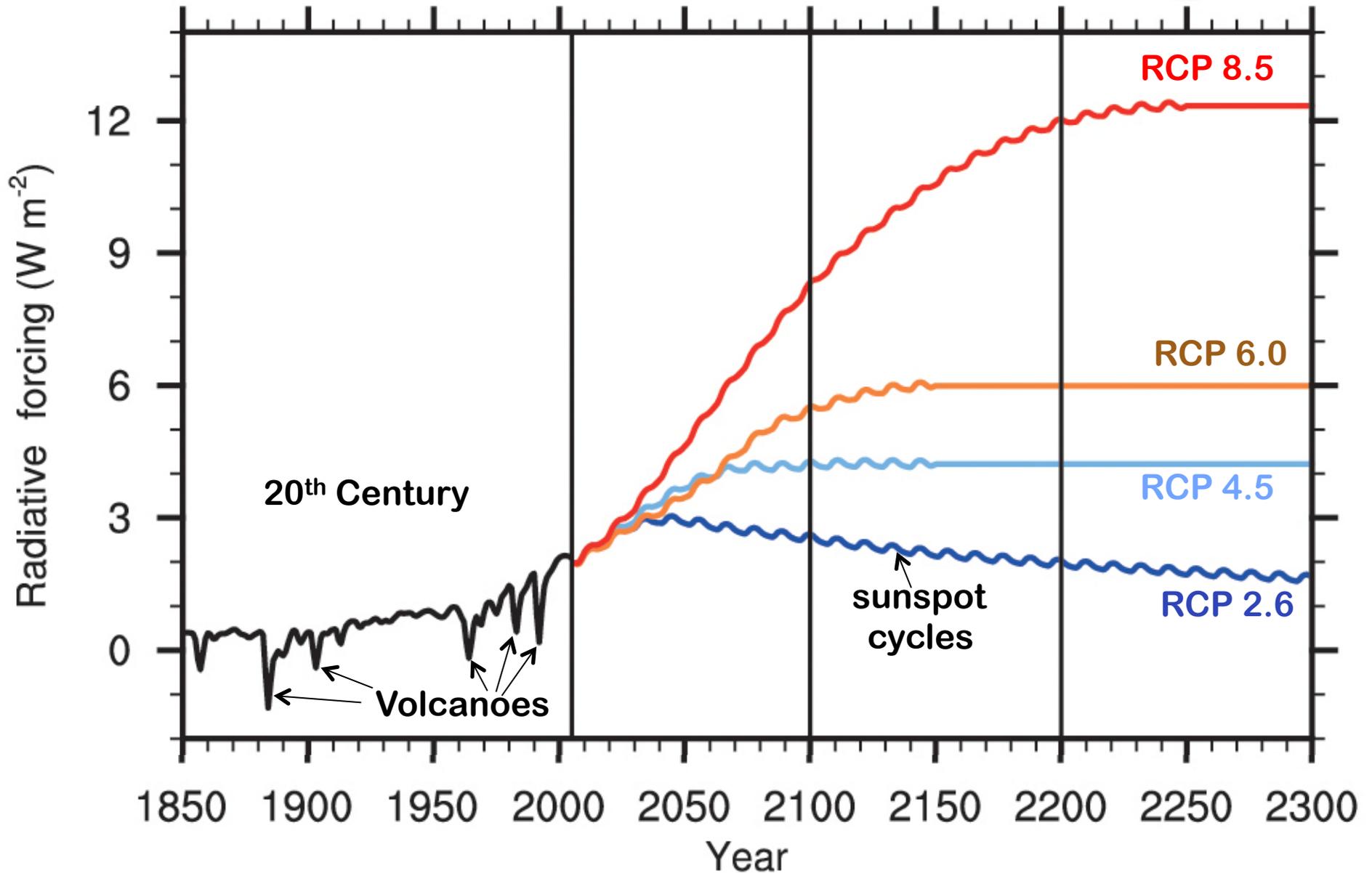
- Population to grow by 40% in 100 years
- Global **economy to grow by 1600%**  
(assumes 2.8% annual GDP growth)

# Scenarios

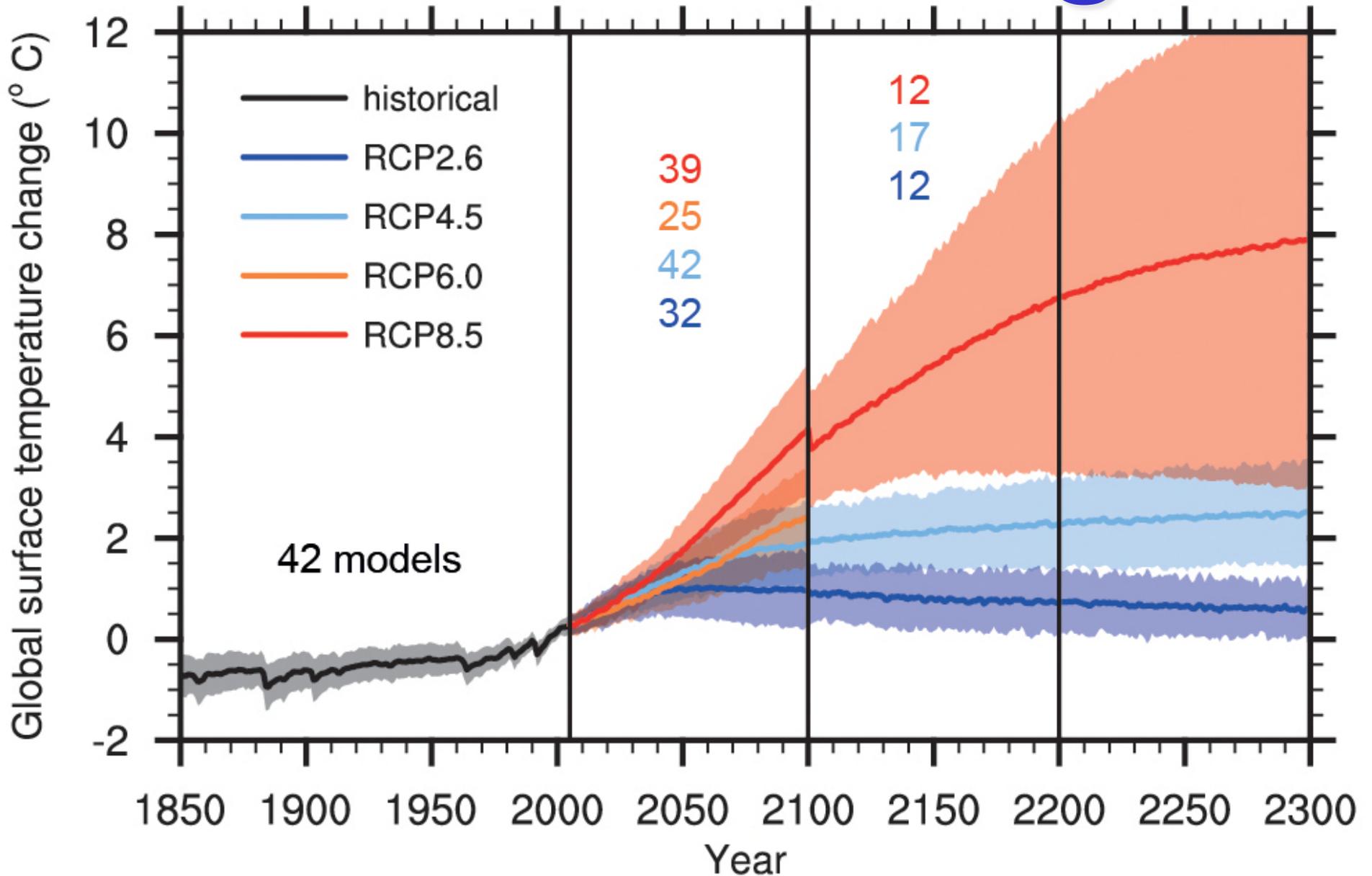
- Not predictions ...  
**“what if”**  
**experiments**
- None are more or less likely
- Depends on  
**economics and policy (politics)**
- Emissions peak in
  - 2020 (RCP2.6)
  - 2040 (RCP4.5)
  - 2070 (RCP6.0)
  - 2150 (RCP8.5)



# Radiative Forcing



# Global Warming

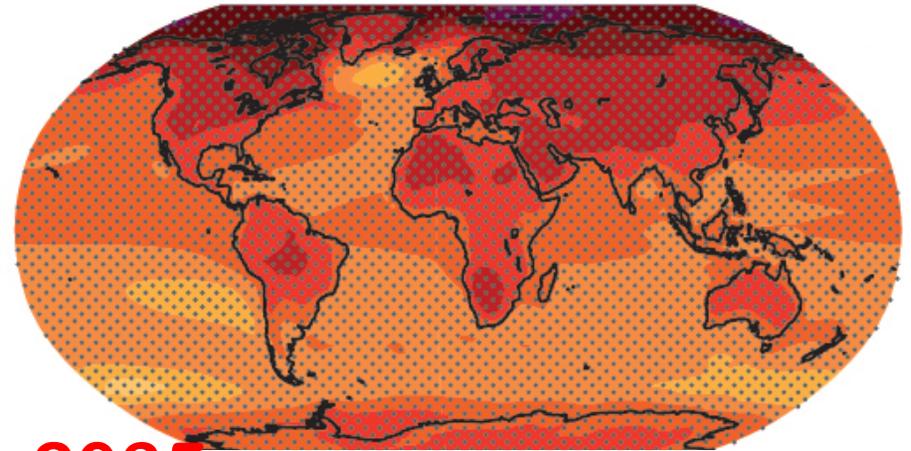


# Maps of Warming

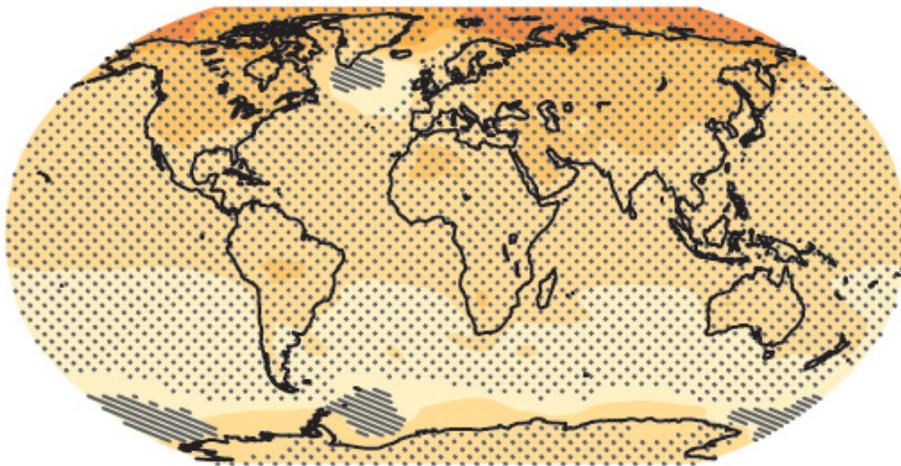
2081 - 2100



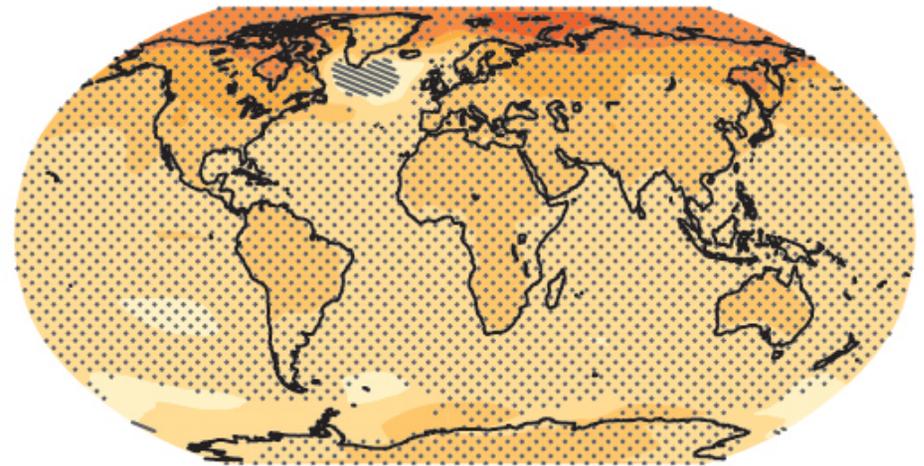
Low Emissions



High Emissions



2016 - 2035

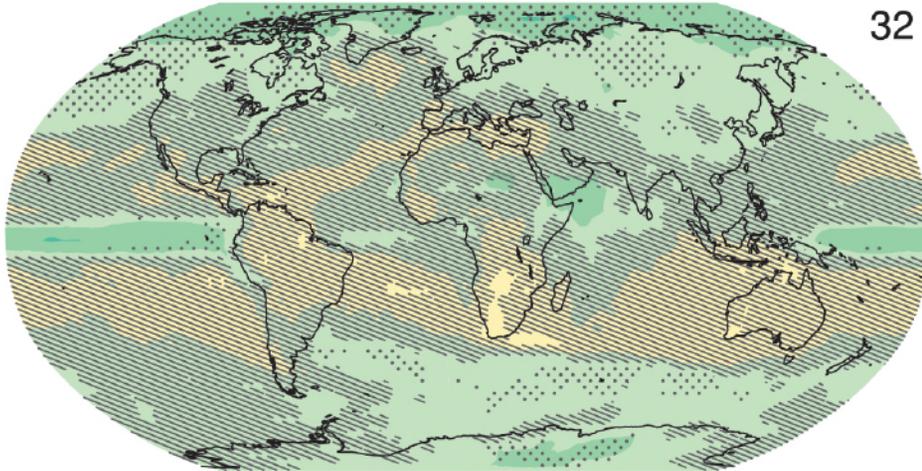


# Precipitation Changes

Annual mean precipitation change (2081-2100)

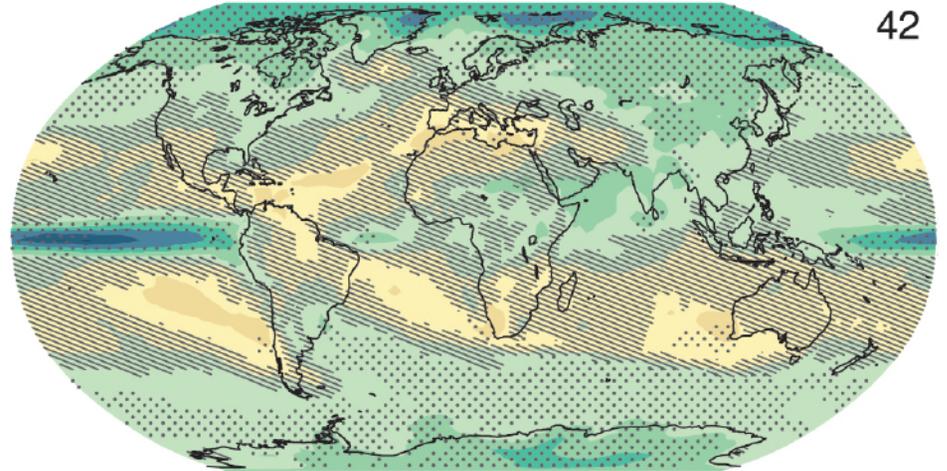
RCP2.6

32



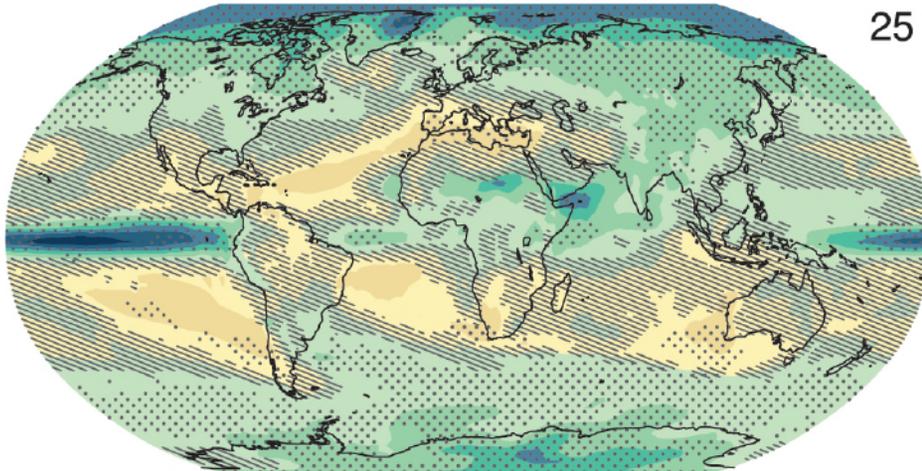
RCP4.5

42



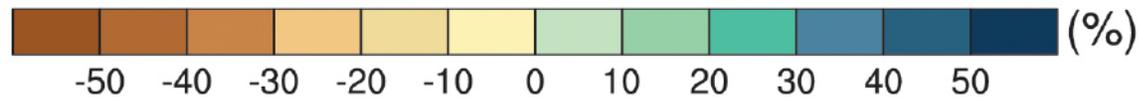
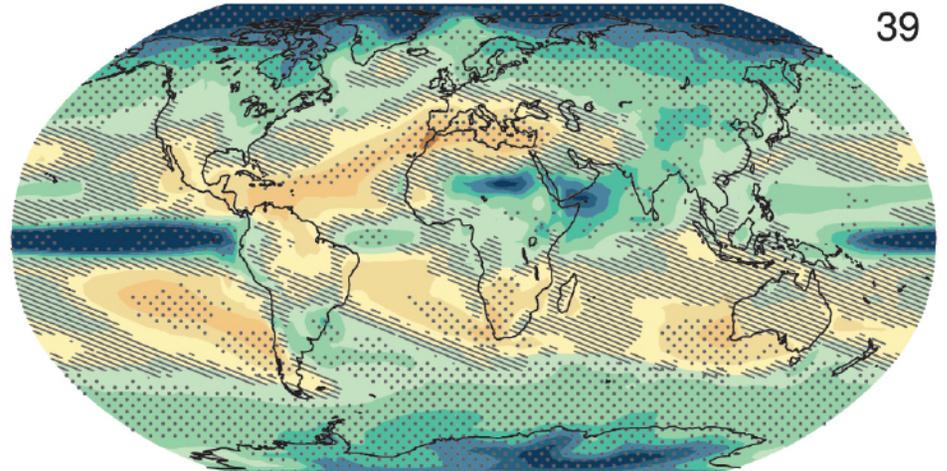
RCP6.0

25

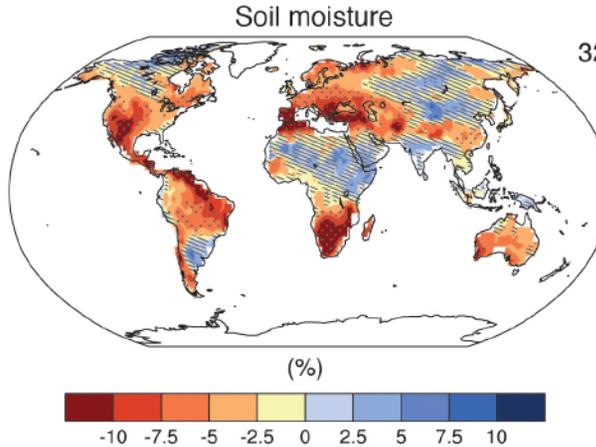
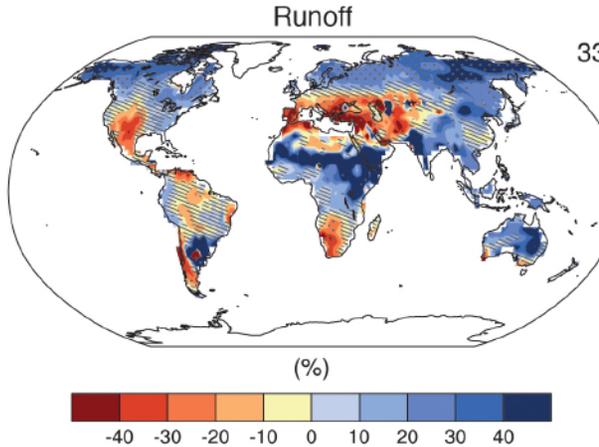
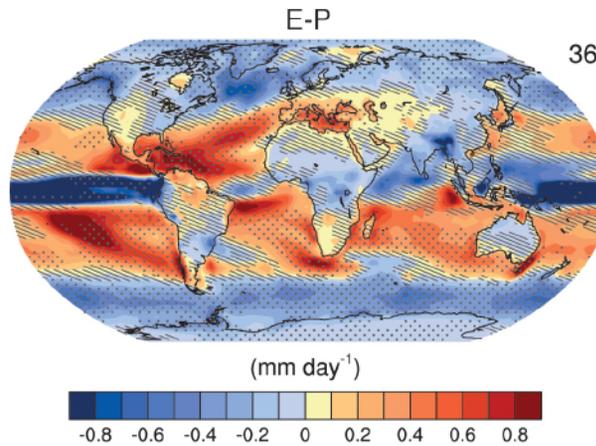
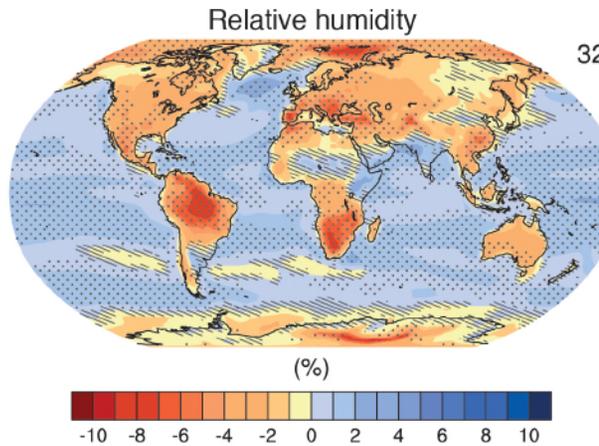
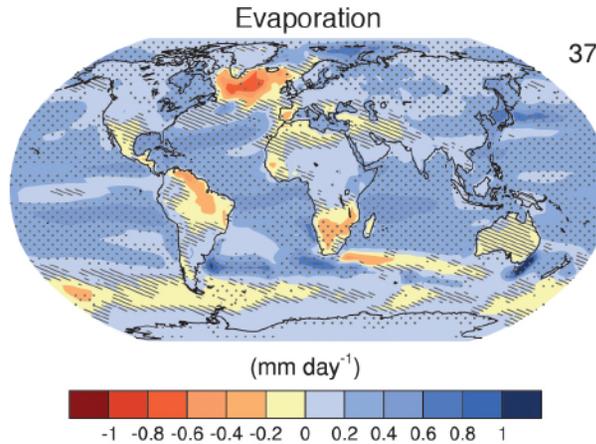
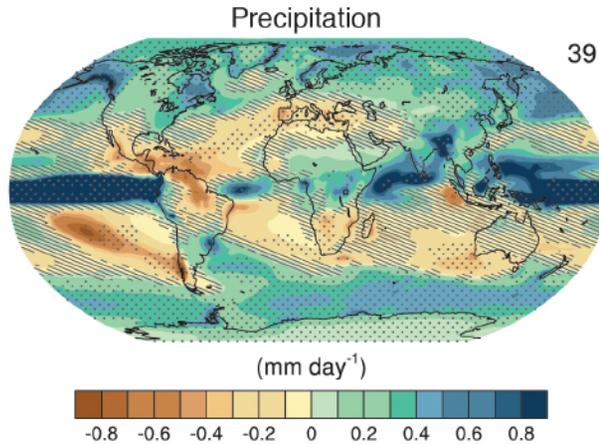


RCP8.5

39



## Annual mean hydrological cycle change (RCP8.5: 2081-2100)

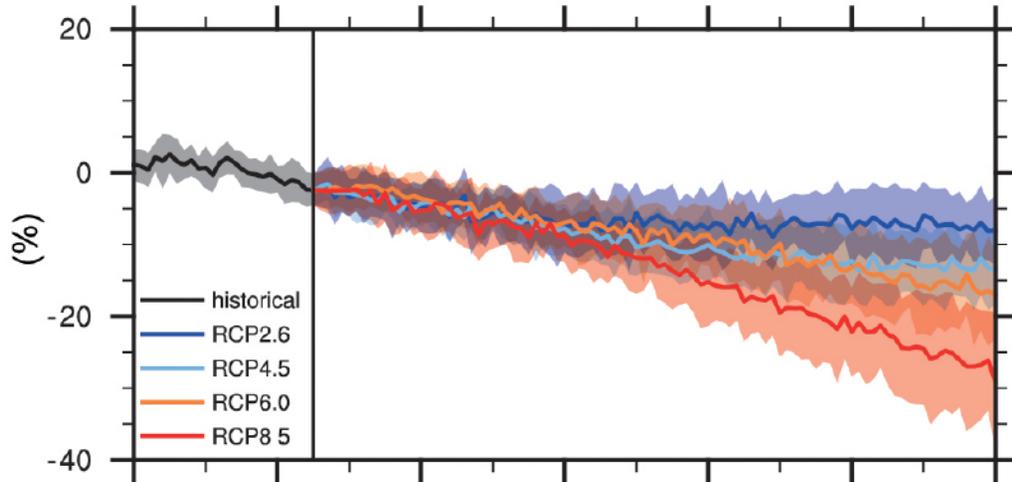


# Changing Hydrologic Cycle

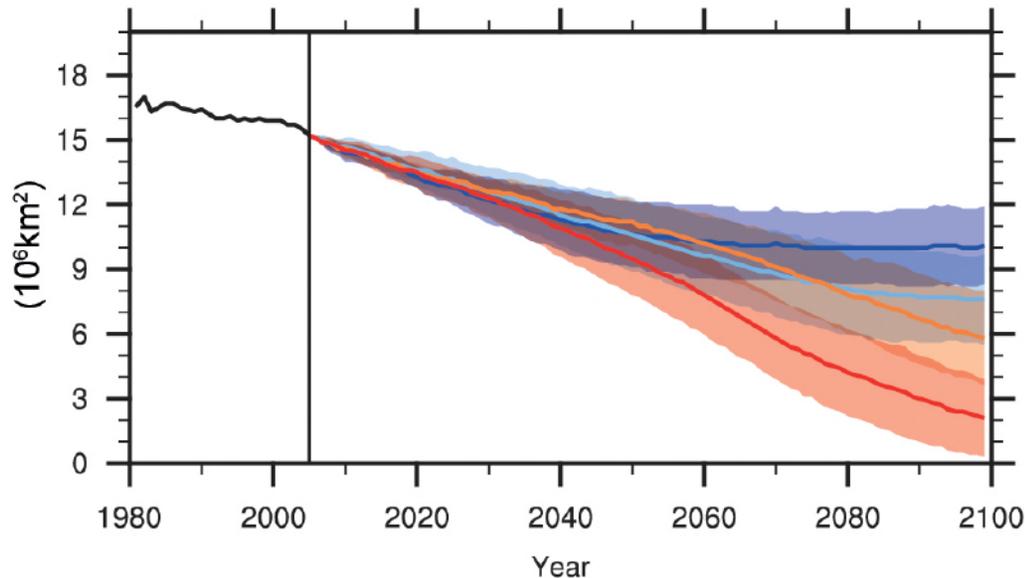
- **Wet places get wetter**
- **Dry places get drier**
- **More evaporation everywhere that there's water to evaporate**
- **Less runoff for reservoirs in many places**

# Snow & Permafrost

Snow cover extent change



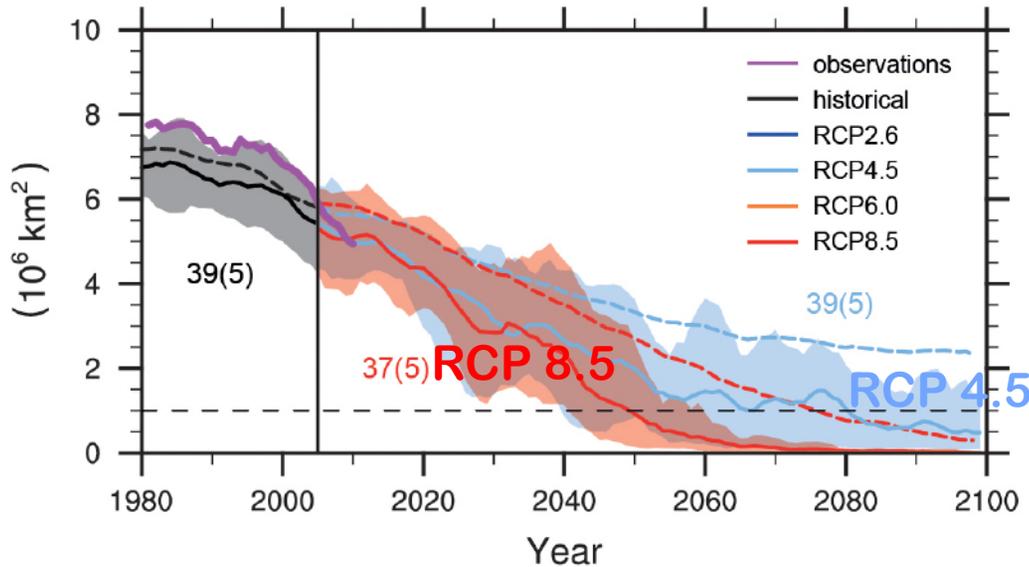
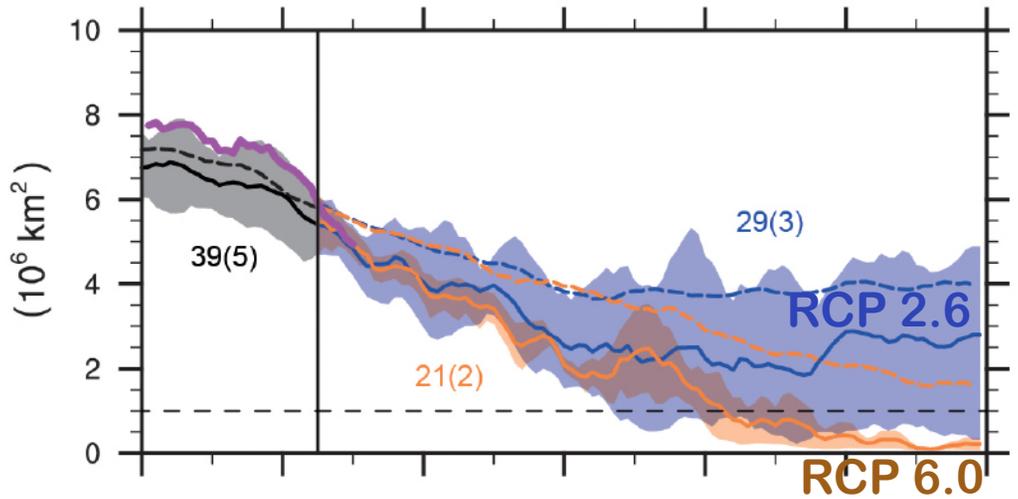
Near surface permafrost area



- Reduced snow and ice under all scenarios
- Reduced snow cover increases absorbed solar radiation (positive albedo feedback)
- Loss of permafrost has special risks ...
  - Drainage
  - Methane and  $\text{CO}_2$

# Sea Ice Changes

NH September sea-ice extent

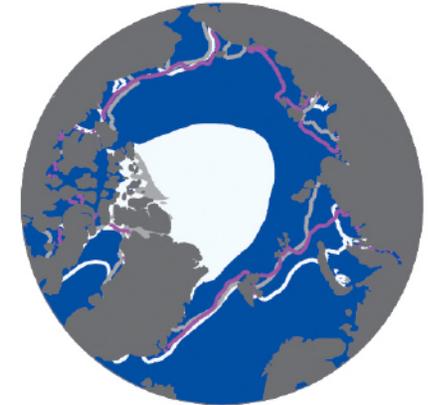


2081–2100

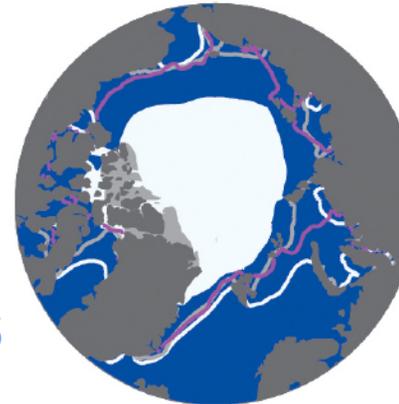
RCP2.6



RCP6.0



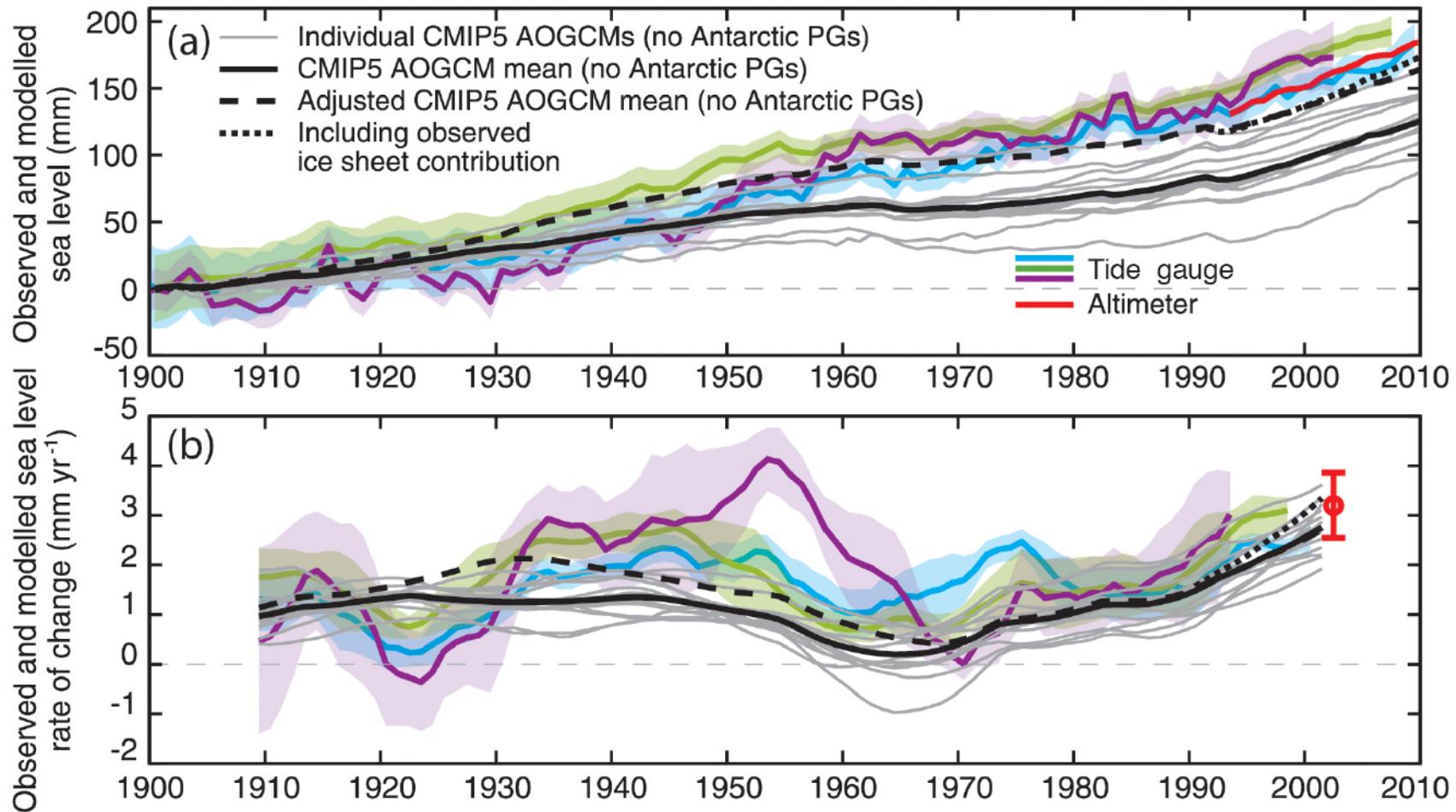
RCP4.5



RCP8.5

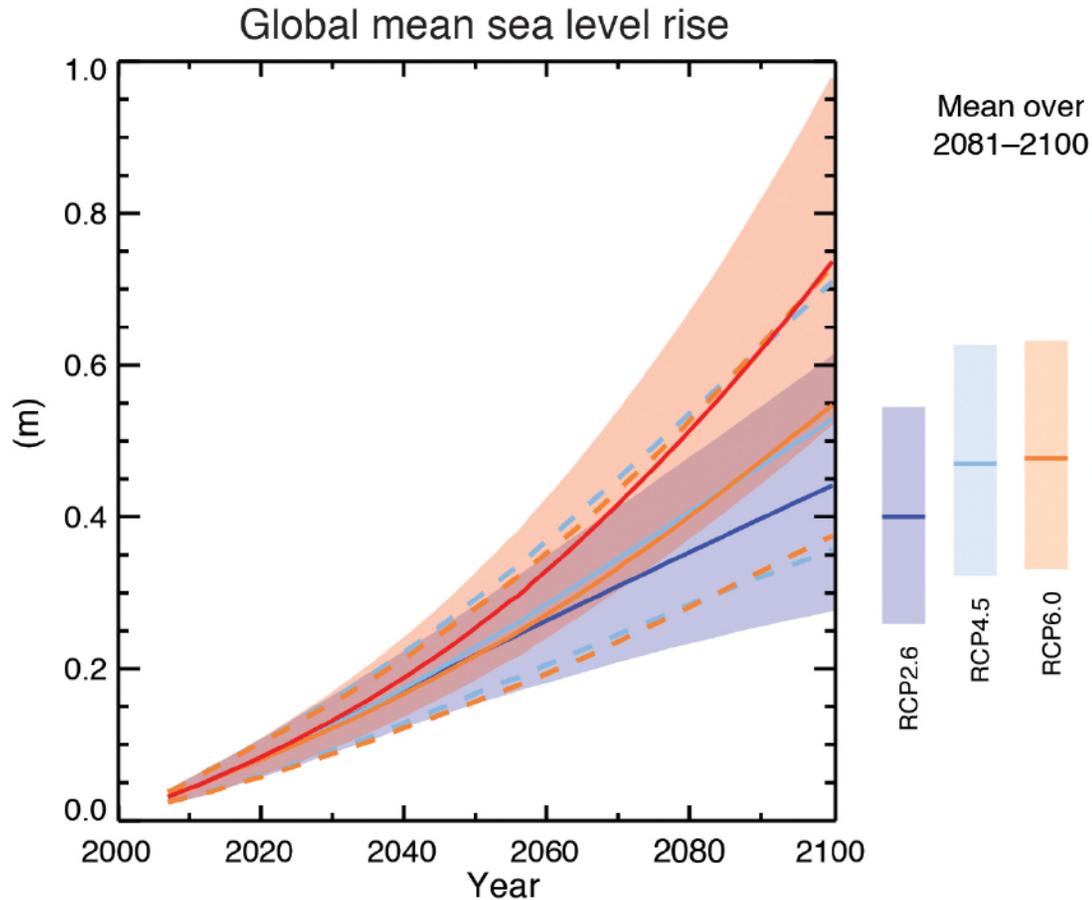


# Evaluation: Sea-Level



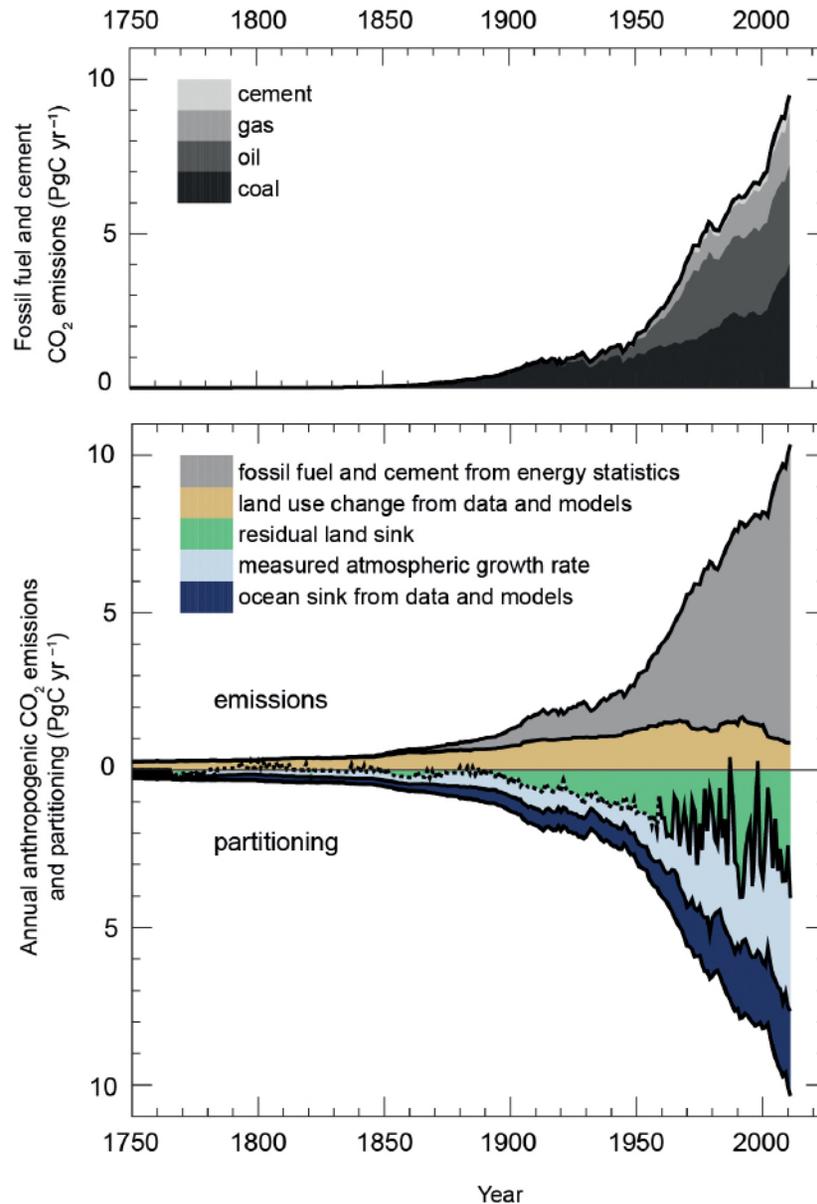
- **Most models underpredict recent sea-level rise**
- **They completely lack ice sheet dynamics**

# Future Sea-Level



- All models show substantial and accelerating rise in sea-level
- Big differences depending on future emissions
- Equilibrium sea level will take many centuries

# Carbon Sources and Sinks

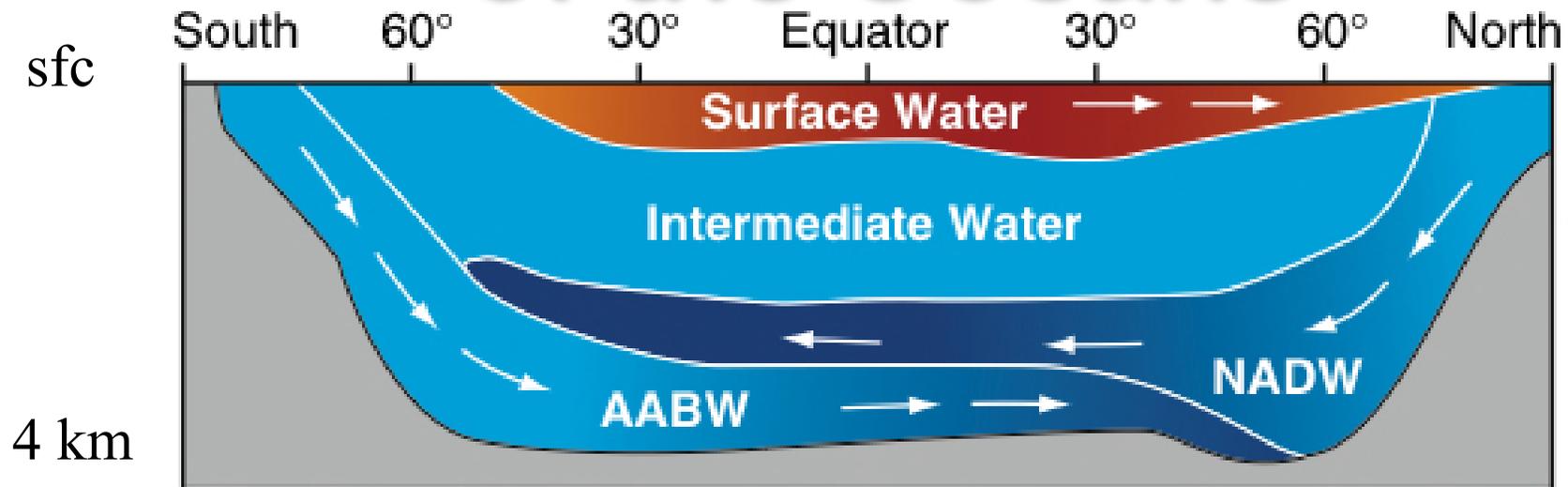


- Half the carbon from fossil fuels goes into the atmosphere
- The other half goes into land and oceans
- The land part is unexpected and unreliable
- Future of carbon sinks is harder to predict than temperatures

# Where Has All the Carbon Gone?

- Into the **oceans**
  - Solubility pump (CO<sub>2</sub> very soluble in cold water, but rates are limited by slow physical mixing)
  - Biological pump (slow “rain” of organic debris)
- Into the **land**
  - CO<sub>2</sub> Fertilization  
(plants eat CO<sub>2</sub> ... is more better?)
  - Nutrient fertilization  
(N-deposition and fertilizers)
  - Land-use change  
(forest regrowth, fire suppression, woody encroachment ... but what about Wal-Marts?)
  - Response to changing climate  
(e.g., Boreal warming)

# Vertical Structure of the Oceans



 Increased nutrients & dissolved CO<sub>2</sub>

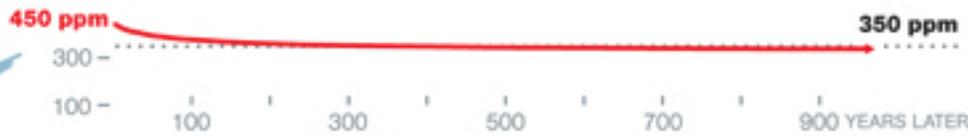
 Warm, low nutrients, & oxygenated

- Warm **buoyant “raft”** floats at surface
- Cold deep water is only “formed” at high latitudes
- Very stable, **hard to mix, takes ~ 1000 years!**
- Icy cold, inky black, most of the ocean **doesn't know we're here yet!**

# Bathtub Drainage

What if we stop emissions completely?

It will take centuries for plants and the ocean to soak up most of the human-made CO<sub>2</sub>. It will take hundreds of millennia for the rest to be removed by rock weathering, which converts CO<sub>2</sub> to carbonate sediments and rocks.



DRAINS FROM TUB

Why would the level stay high for so long?

Plants and soil absorb CO<sub>2</sub> quickly, but that reservoir fills up fast.

The deep ocean is bigger, but access is slow; CO<sub>2</sub>-laden surface water sinks at only two places near the Poles.

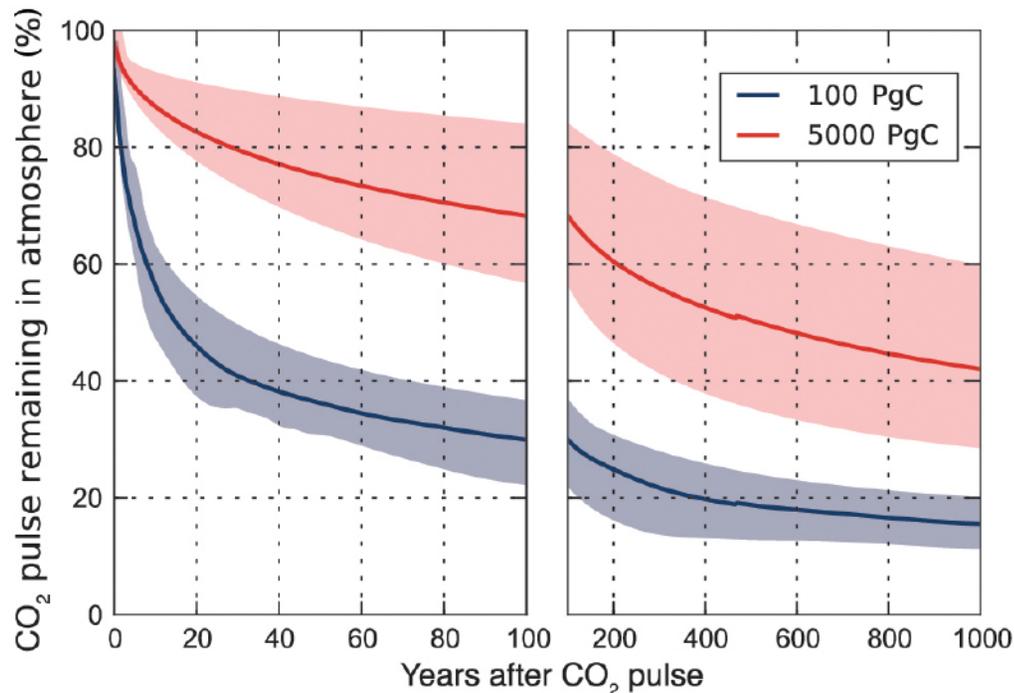
Carbonate sediments and rocks are far bigger and slower still; they form at sea from elements weathered off rocks on land.

PLANTS & SOIL

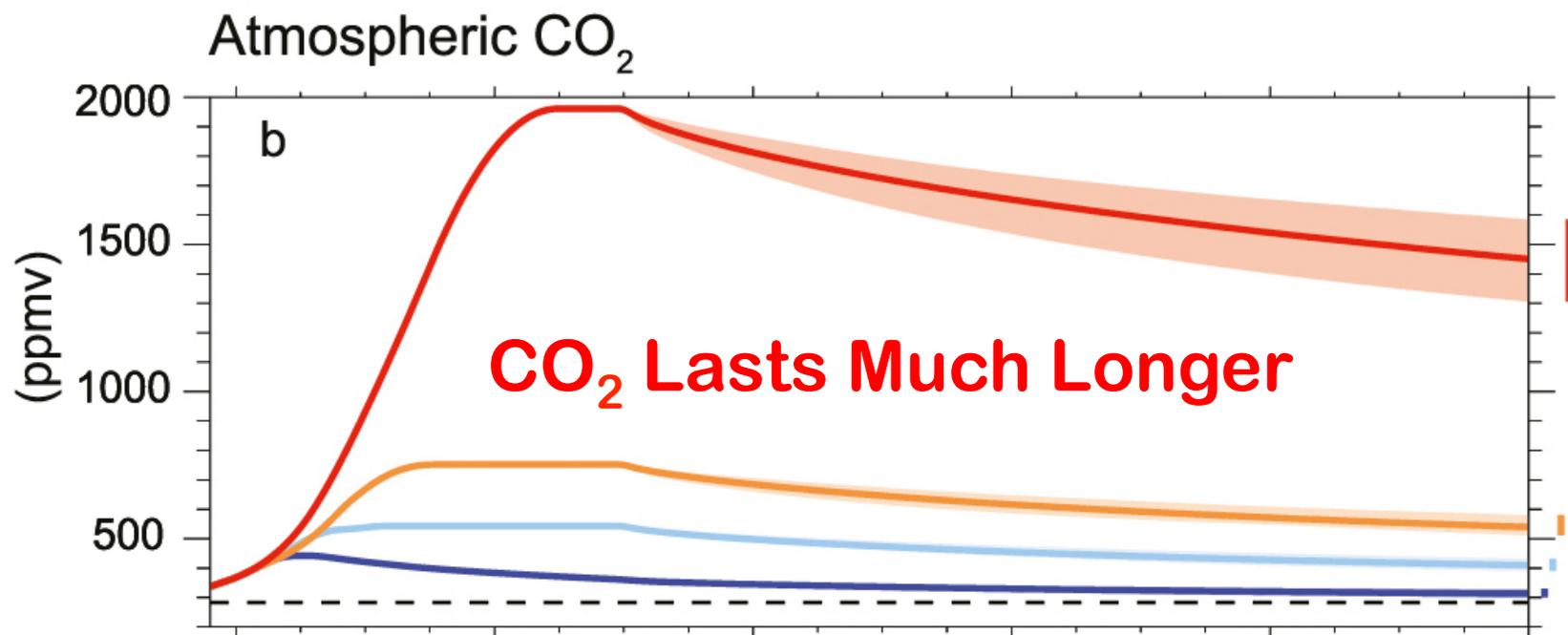
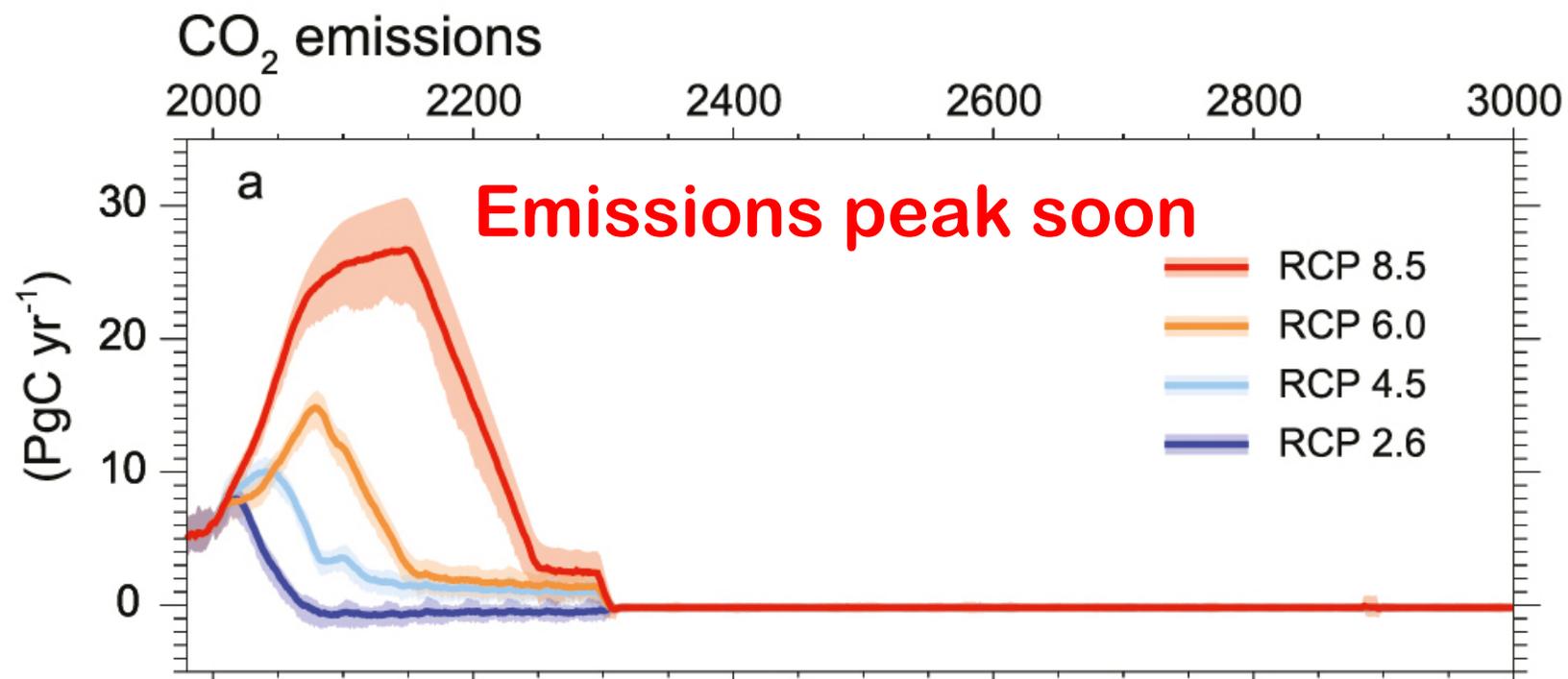
DEEP OCEAN

SEDIMENTS & ROCKS

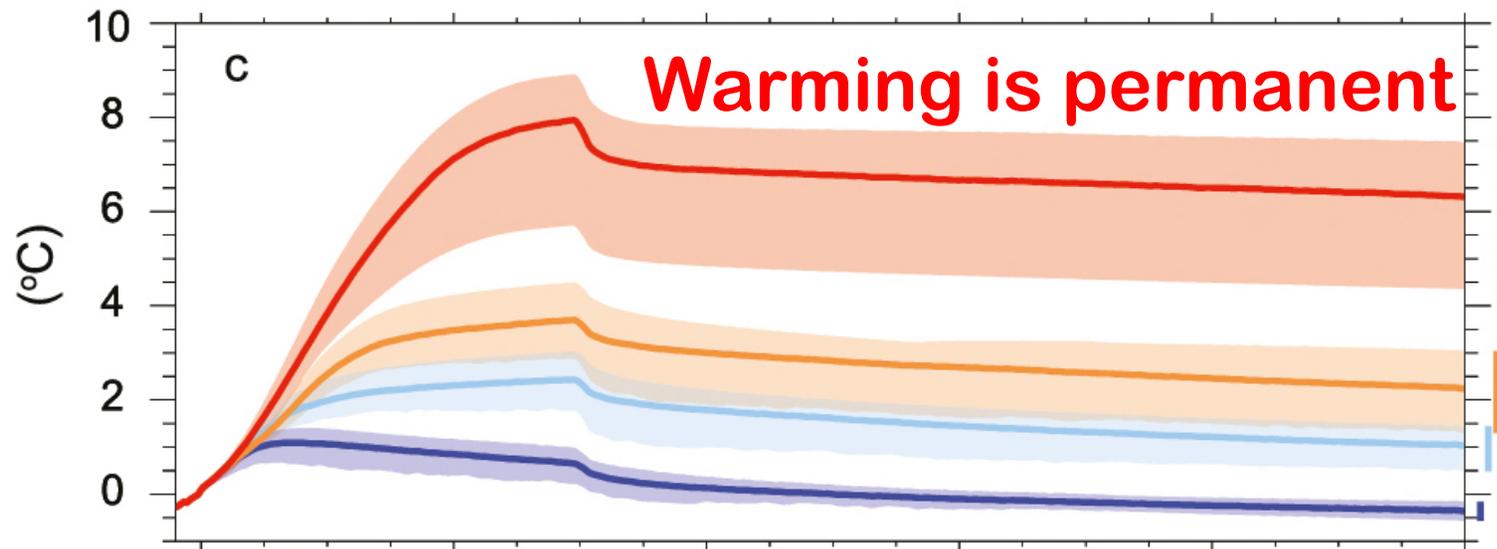
# The Long Tail



- Fossil CO<sub>2</sub> dissolves into the oceans
- Chemistry limits the amount the oceans can hold
- **Mixing with deep oceans is very slow**
- Removal of CO<sub>2</sub> depends on how much we add to atmosphere
- For a big pulse, **40% is still in the air after 1000 years**



## Surface air temperature change



## Ocean thermal expansion

